

Mesoscale



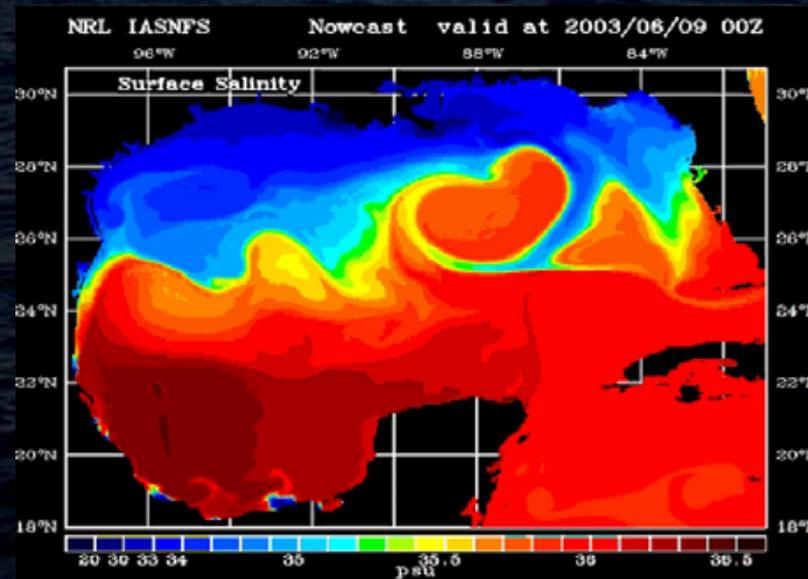
Climate change

Synoptic dynamics

Subsurface inference

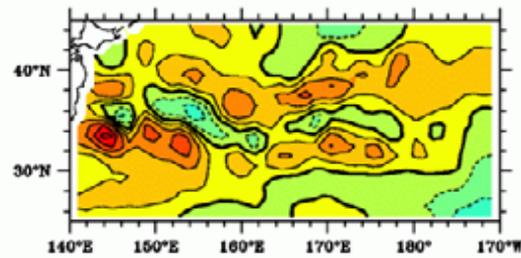
Dynamic accuracy

Future directions for systems
and applications

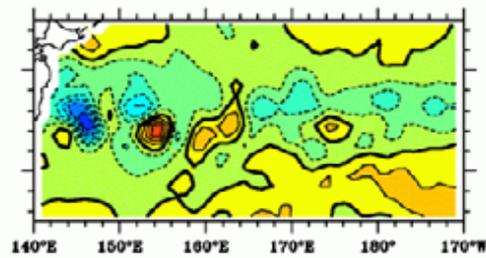


T/P yearly SSHA fields

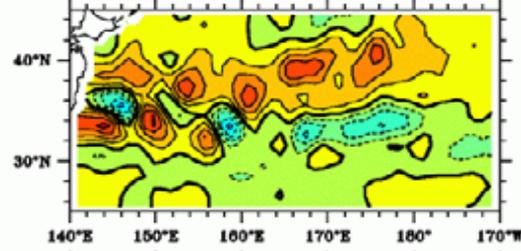
1993



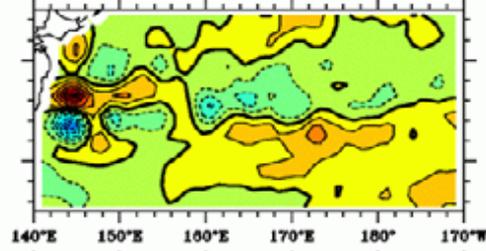
1998



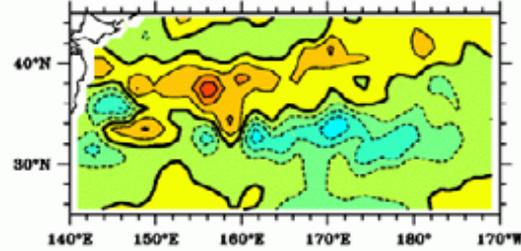
1994



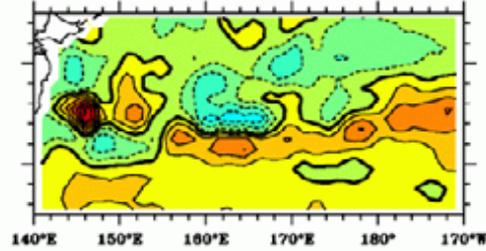
1999



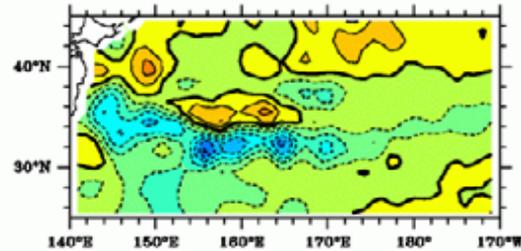
1995



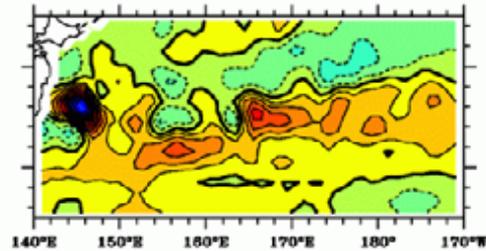
2000



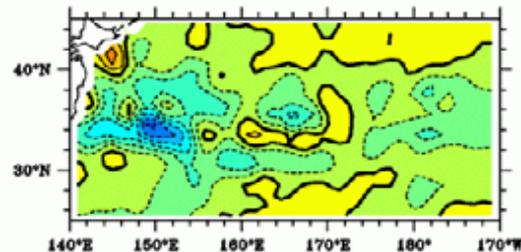
1996



2001

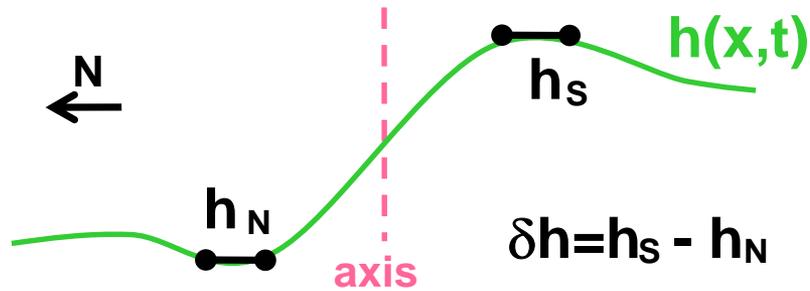
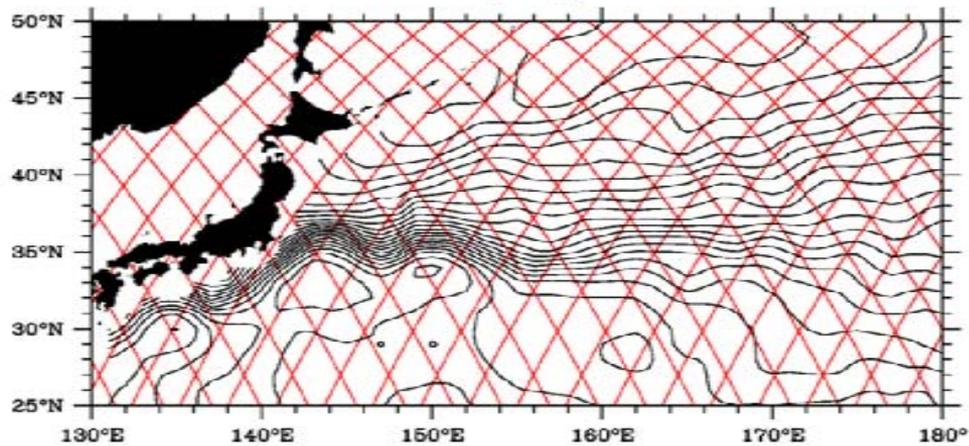


1997

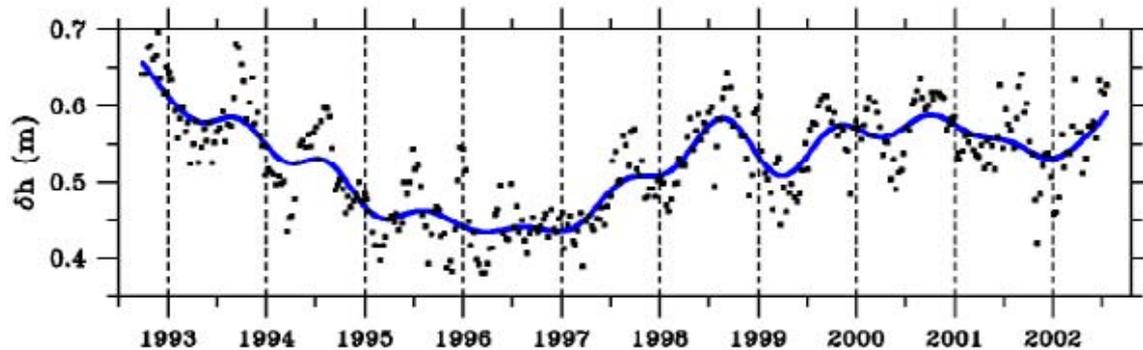


Qui et al., Dec. 2003

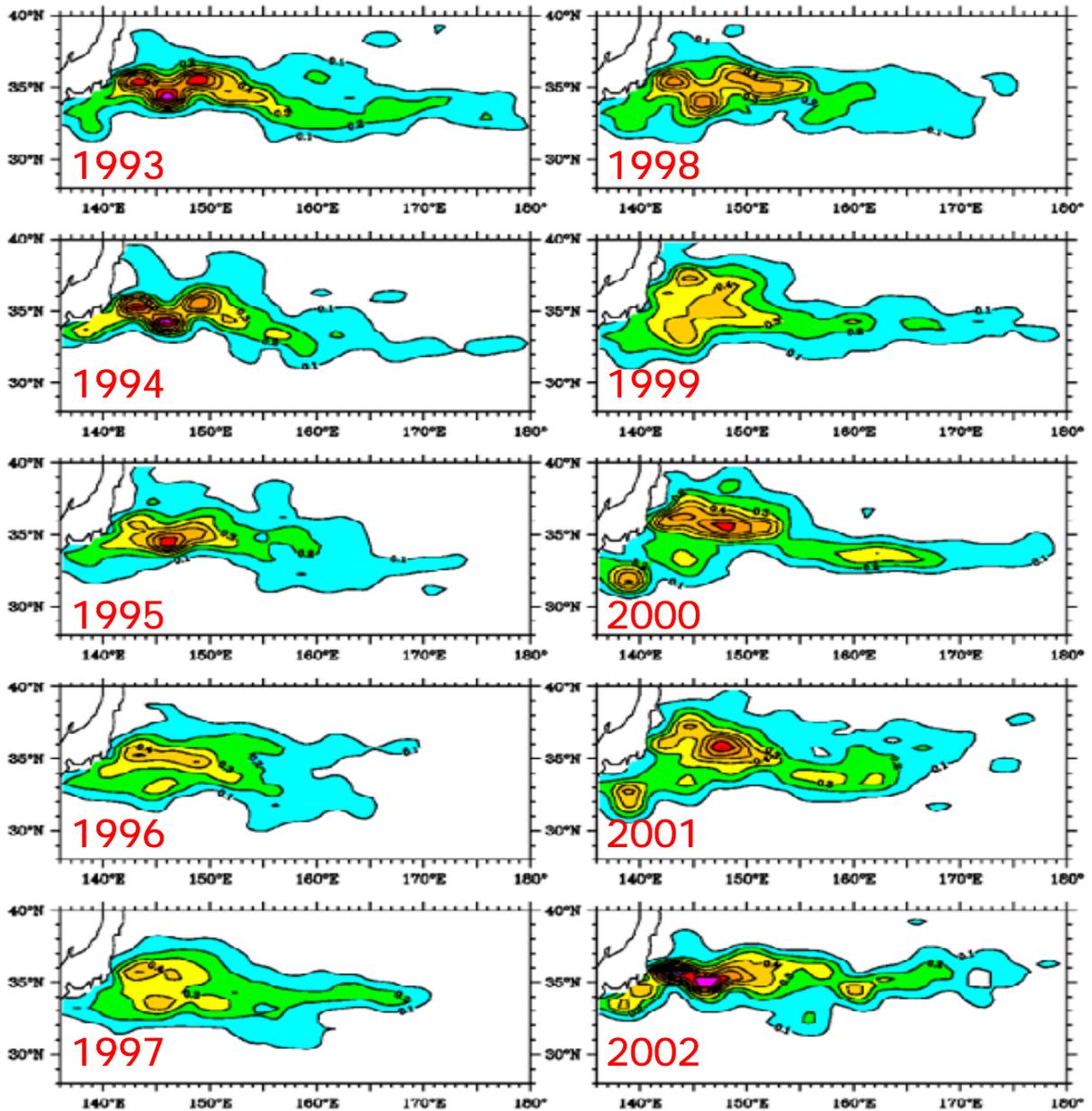
Mean SSH vs. T/P ground tracks



$\langle \delta h \rangle$ across KE axis averaged in 142°E–180°



Yearly kinetic energy maps



Mesoscale

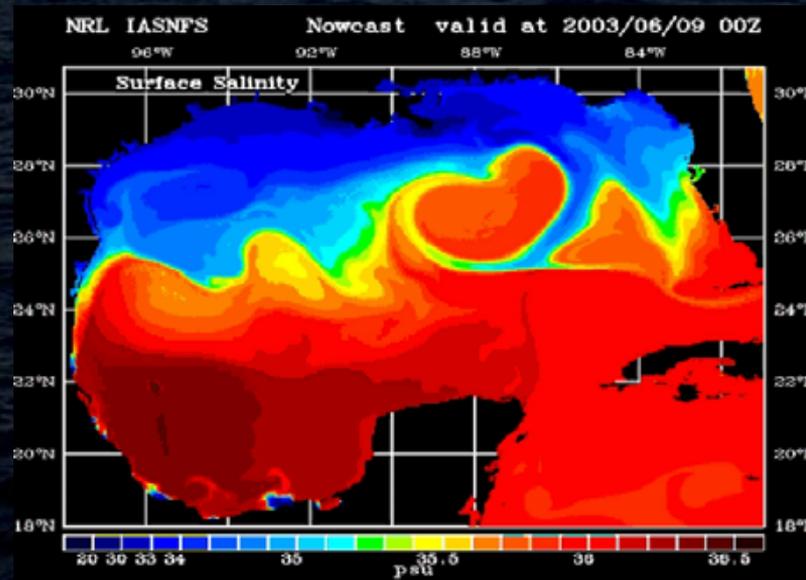
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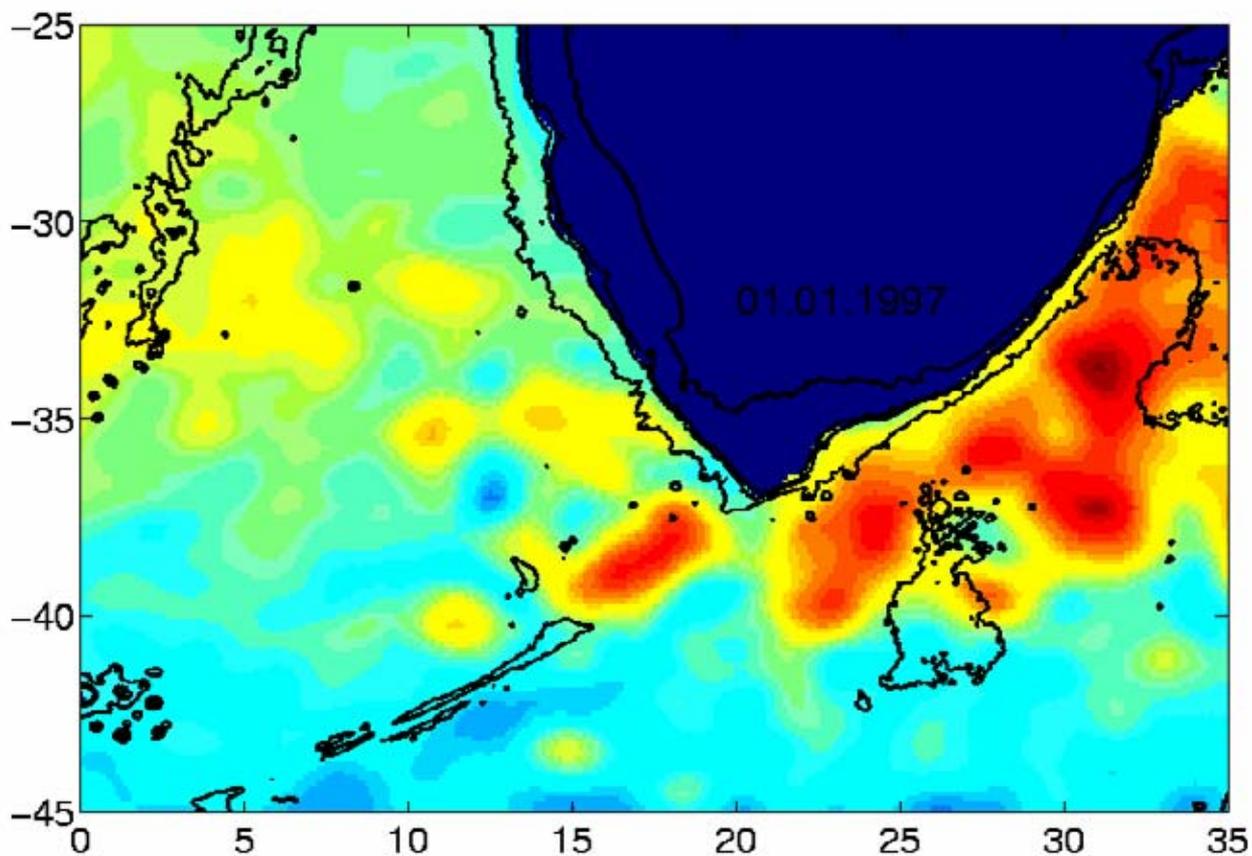
Future potential systems and applications





Agulhas

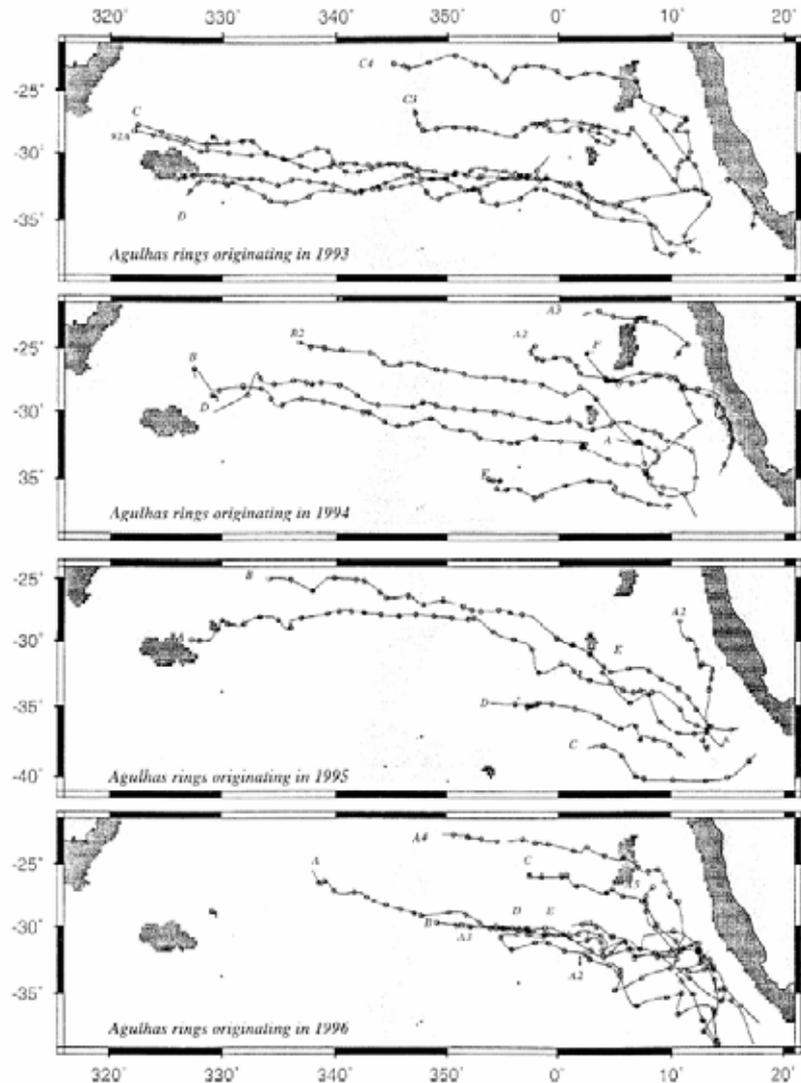
Boebel, O. and C. Barron, 2003.





Observing Mesoscale

Mesoscale Agulhas ring position observations.



1993

1994

1995

1996

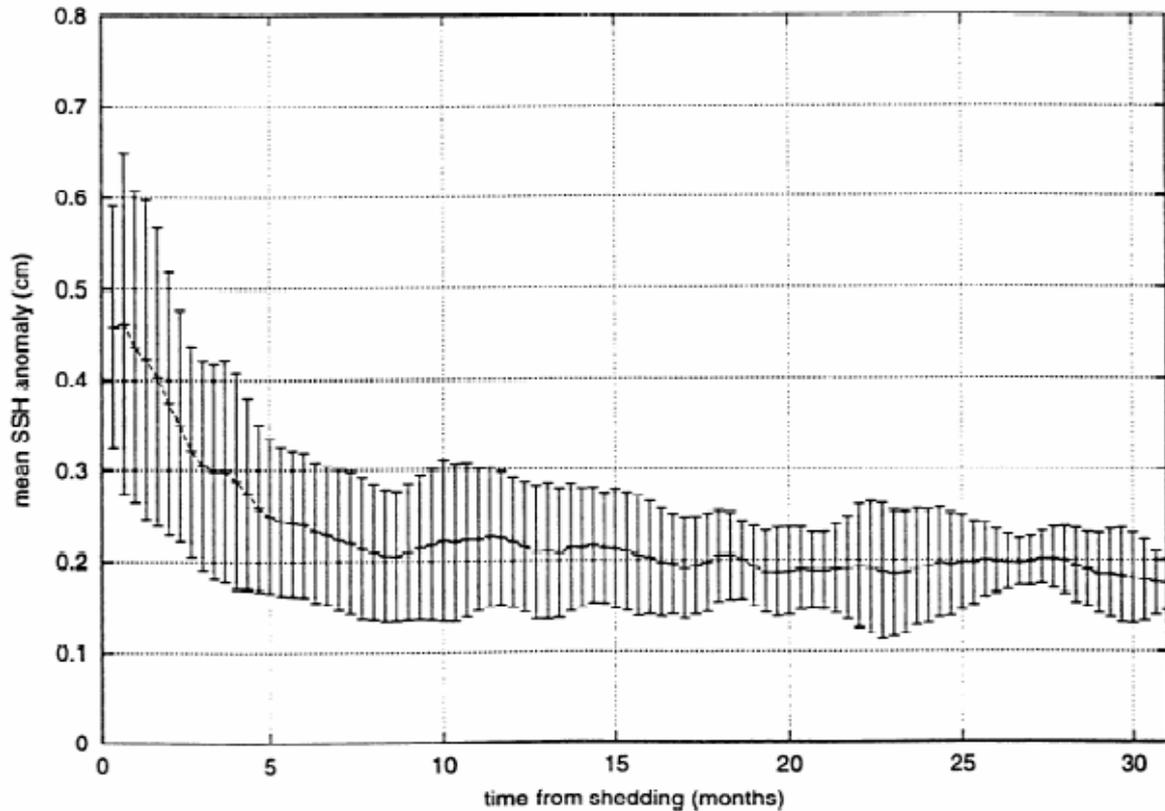
Schouten et al., JGR 2000.



Mesoscale Characteristics

Decay rate over time
(following each
individual eddy).

Ring energy from theory
is proposed to decrease
rapidly during first 5
months due to Rossby
wave radiation
(Beismann et al., 1999),
and then stabilize
afterward. Observations
confirm this theory.



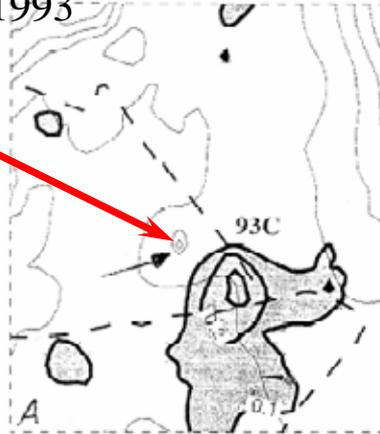
Schouten et al., JGR 2000.



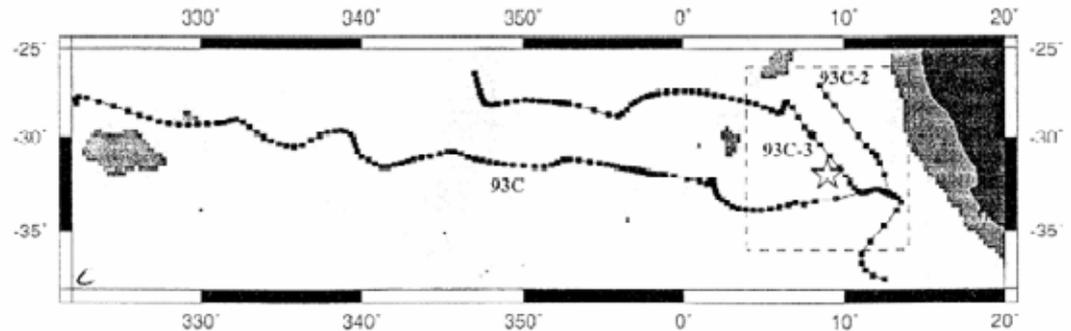
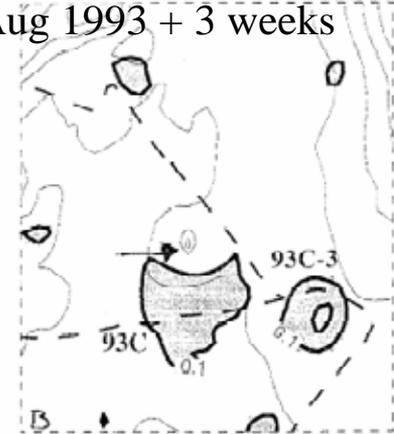
Mesoscale Evolution

Eddy splitting over
Vema Seamount.

Aug 1993



Aug 1993 + 3 weeks



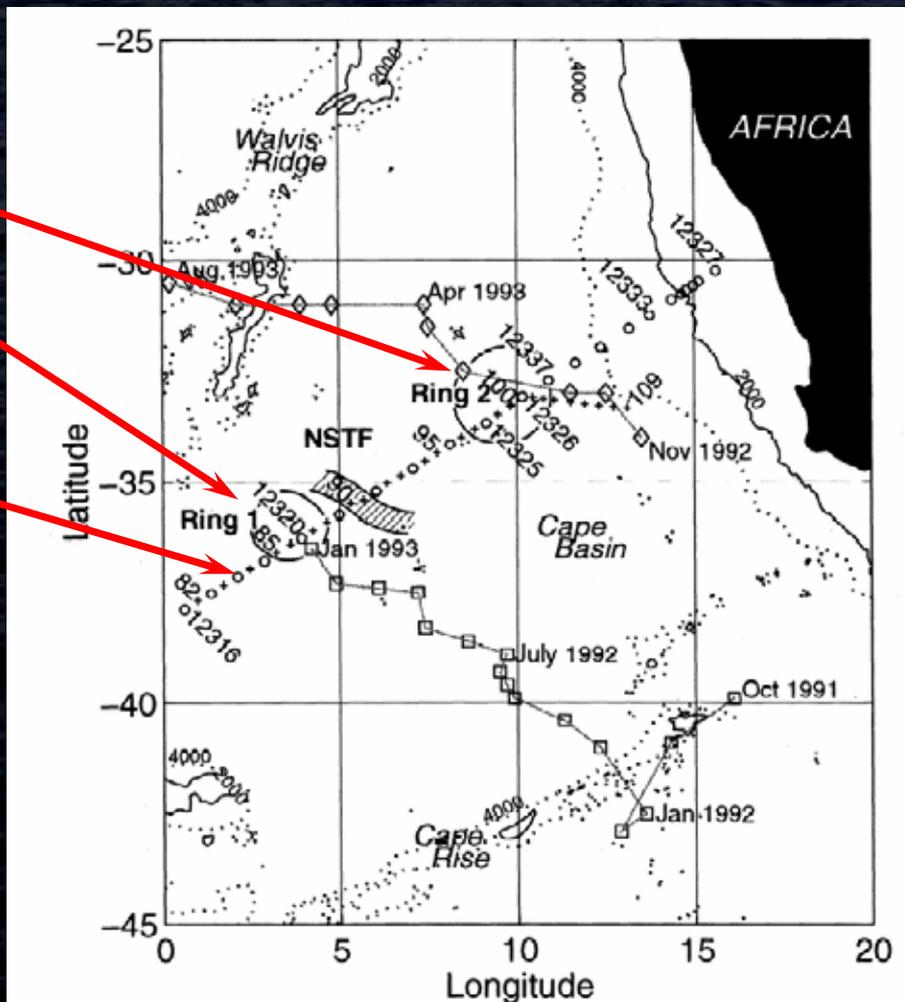
Schouten et al., JGR 2000.



Mesoscale Observations

Agulhas rings

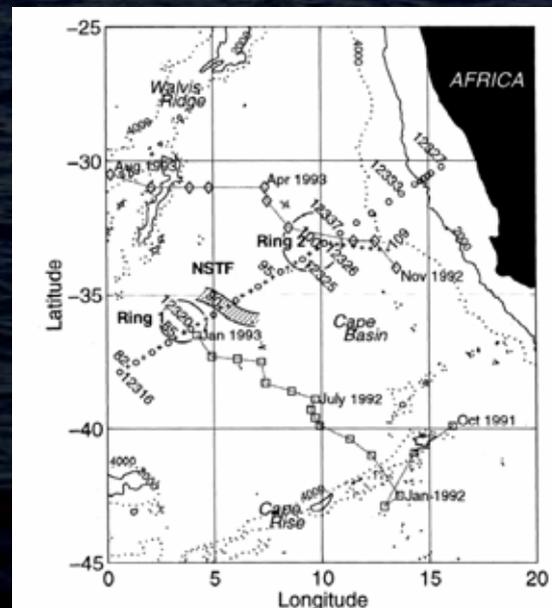
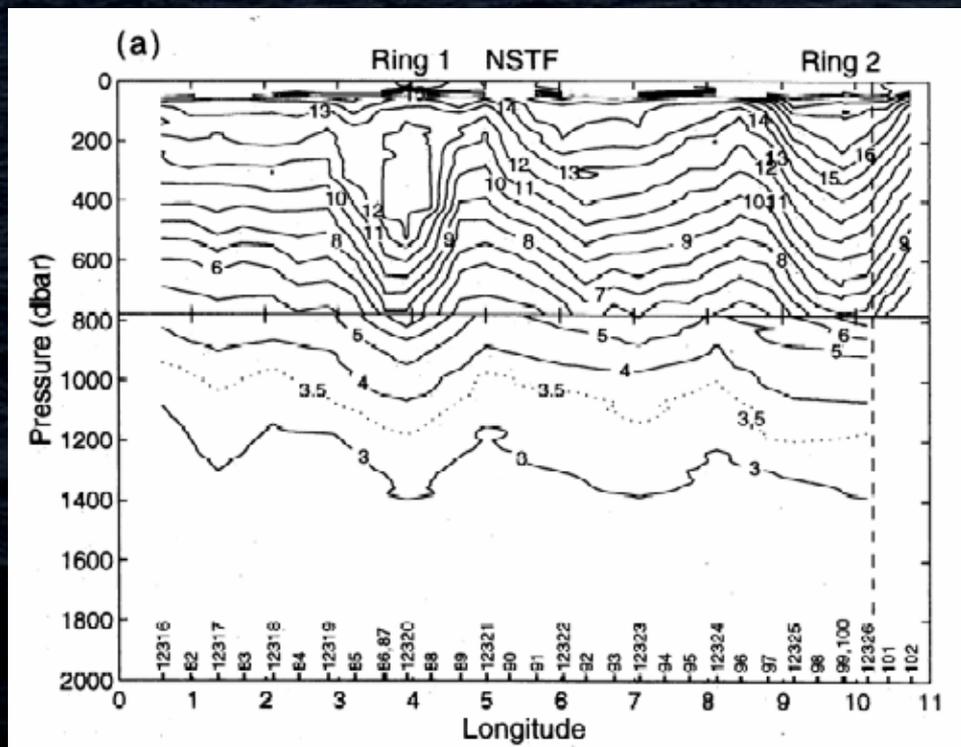
Ship transect



McDonaugh, et al., JGR 1999.



Mesoscale Observations



Potential temperature (°C)

Based on observed T&S profiles and geostrophic velocity, heat and salinity fluxes are estimated.

McDonaugh, et al., JGR 1999.



Mesoscale Observations

Observed propagation speeds of Rings 1 and 2:
2.8 and 6.5 cm/s

Theoretical self-induced speeds due to β effect:

Table 3. Limits of Westward Velocities Induced by the β Effect for Rings 1 and 2

	Nof [1981]		Cushman-Roisin <i>et al.</i> [1990]	
	$\frac{1}{3}\beta R_D^2$	$\frac{2}{3}\beta R_D^2$	$\beta g' (h_w + dh) / 2f^2$	$\beta g' h_w / f^2$
Ring 1	0.4	0.7	-	0.8
Ring 2	0.4	0.9	1.4	1.5

Velocities are in centimeters per second. R_D is radius of deformation, g' is acceleration of gravity; h_w is depth of the 10°C isotherm outside of the ring; and f is the Coriolis parameter.

Therefore advection is a significant effect on eddy propagation.

Mesoscale

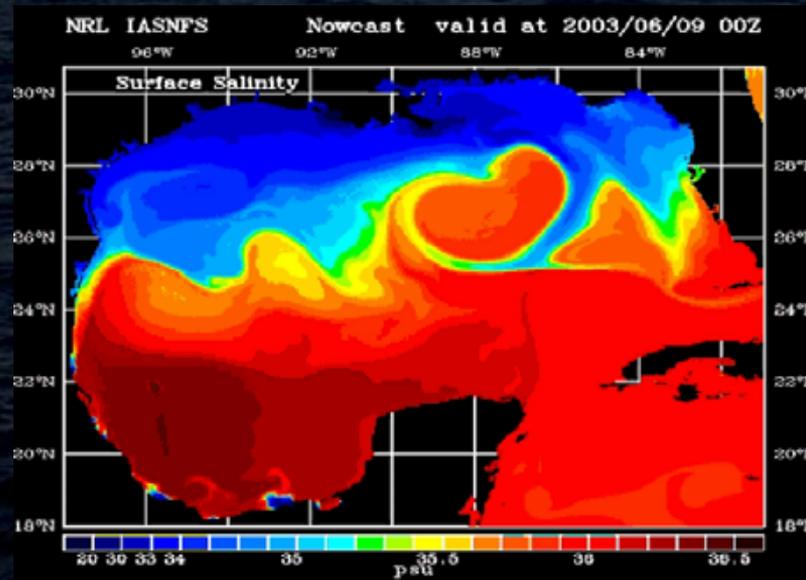
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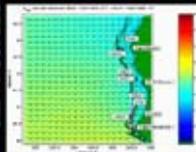
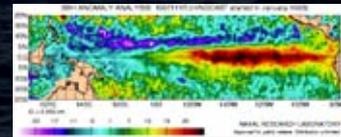
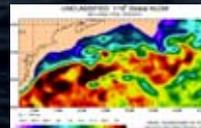
Future potential systems and applications





Purpose for Ocean Assimilation

- **Non-deterministic features require continuous observation. Predictive time scales are order of 1-2 weeks for atmosphere and 15-30 days for ocean.**
 - Mesoscale eddies
 - El Nino
- **Correction of forcing fields**
 - Wind-driven circulation
 - Wind-driven wave field
 - Heat-driven steric variability
 - Boundary conditions for high resolution local models (tides for example)
- **Correction of inaccurate dynamics**
 - Tides errors due to bathymetry inaccuracies (part of dynamical equations)
 - Vertical turbulence closure inaccuracies leads to errors in vertical density distribution
 - Assimilation should not be a replacement for accurate dynamics
 - A system **MUST** represent the dynamics contained in the input data





Need for Observations

An accurate numerical model should reproduce realistic statistics without assimilation

- Mean flow, position of currents
- Eddy variability
- Eddy statistics (size, speed, energy, ...)

Without assimilation, modeled mesoscale features do not match synoptic reality.

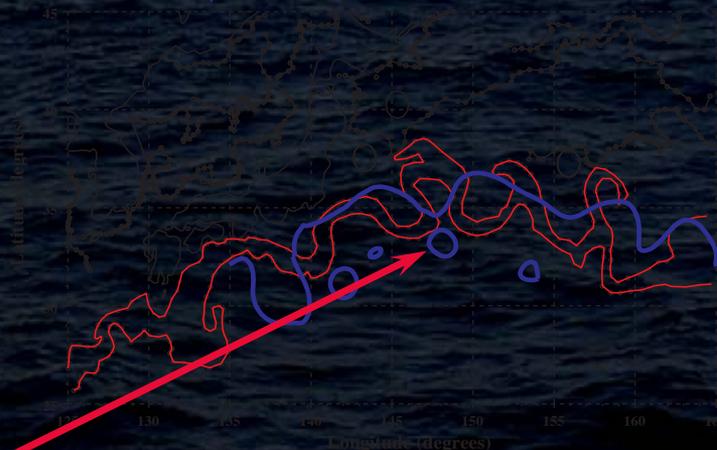
With assimilation, modeled mesoscale features match synoptic reality.

..... Observed IR position

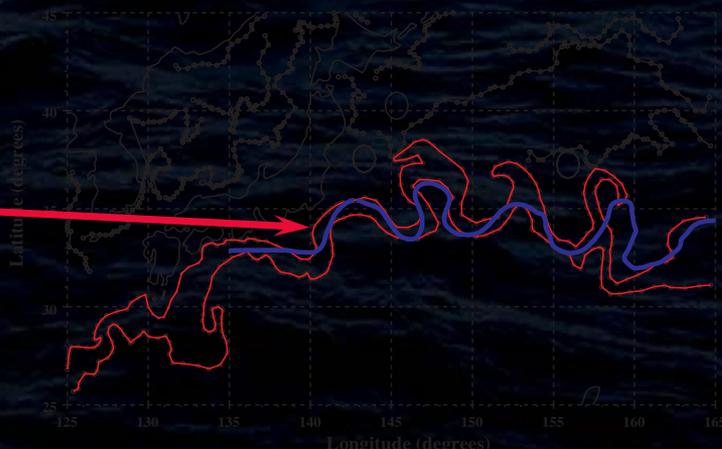
——— Model position

June 25, 1993

NO ASSIMILATION



ERS-1 and TOPEX ASSIMILATION



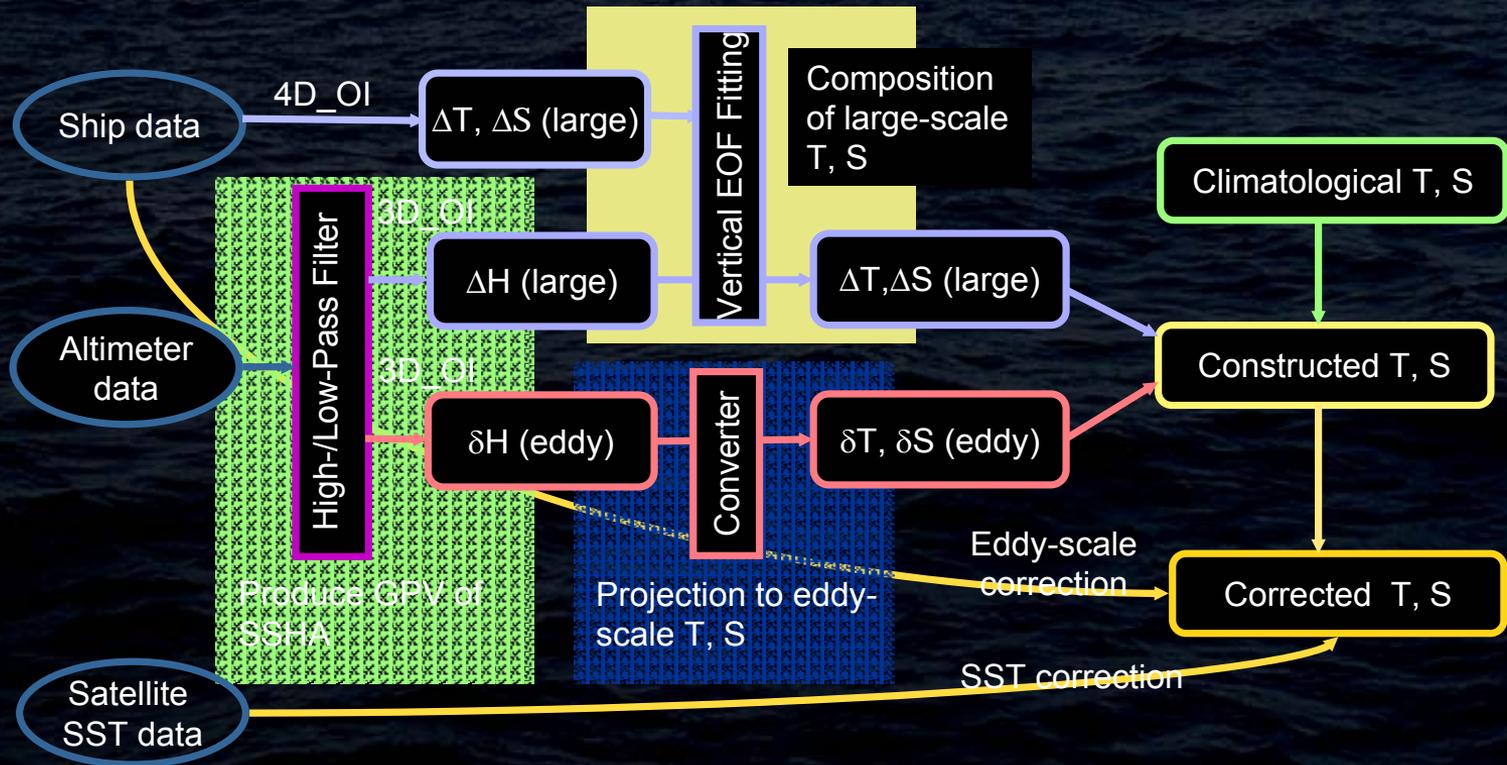
——— IR frontal position

——— model frontal position

Smedstad et al., 1998



Japan Met Agency Analysis Cycle (Kuragano)



4D_OI: space(3D)-time(1D) optimum interpolation. 3D_OI: space(2D for horizon)-time(1D). Converter: statistical regression method. Vertical EOF fitting: linear composite of EOF modes to fit OI analysis.



Indirect Observations (Surface Connection to Subsurface)

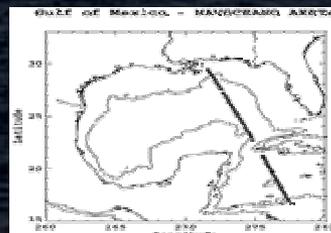
The deep ocean mesoscale variability is due primarily to displacement of waters of different density.

It is possible (through historical in situ observations) to relate surface height and temperature to subsurface temperature and salinity (and thus density).

Assimilation of only surface data leads to a very slow convergence of the model to observations.

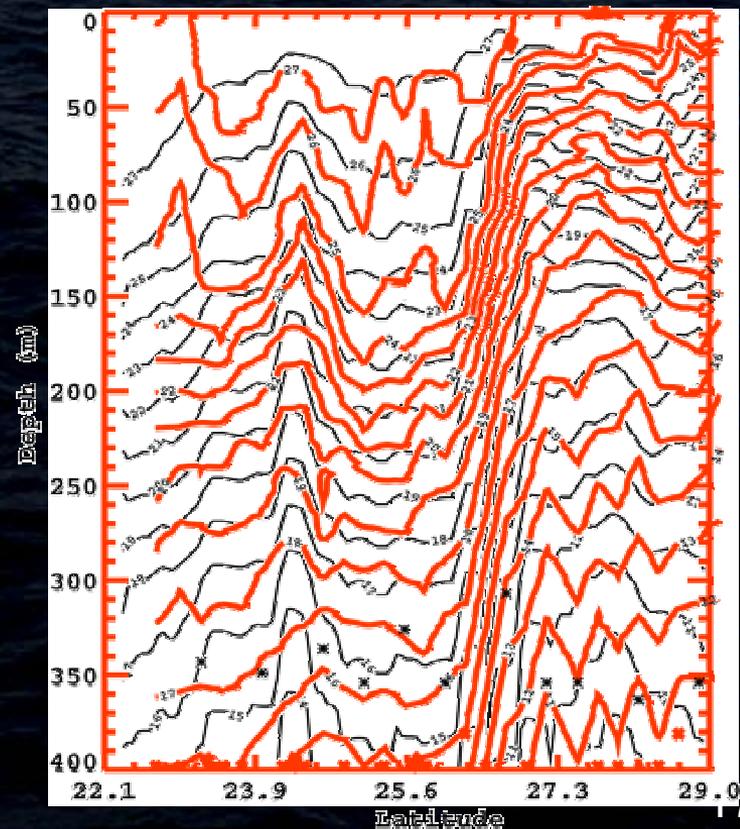
This “synthetic observation” is valuable information to extend the surface observations below the surface rapidly in the assimilation system.

Another way to view this is that we are applying all our knowledge to the problem including historical observations and statistics.



AXBT survey during May 1999

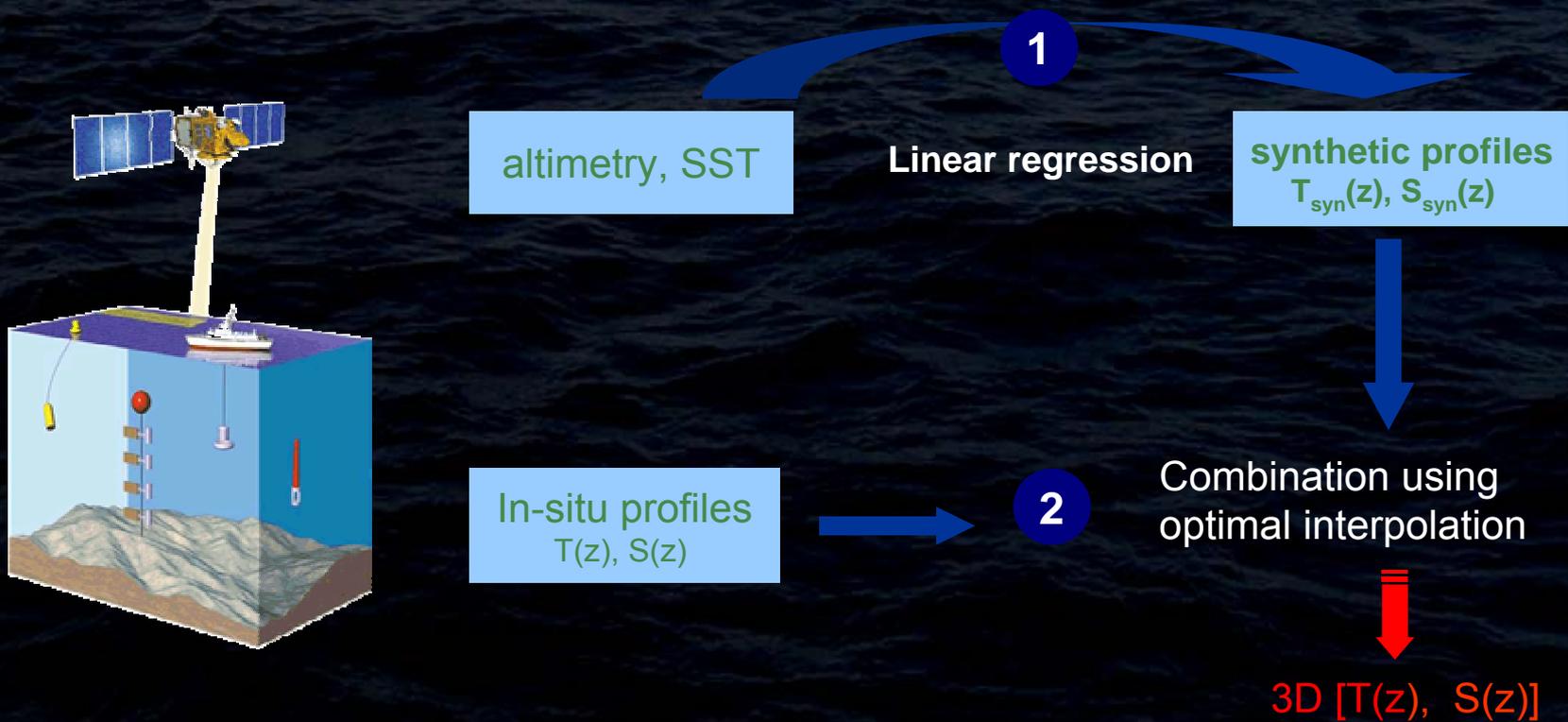
AXBT vs. MODAS/GFO





Data Combination (Le Traon)

- 1 vertical projection of altimetry + SST
- 2 Combination of the synthetic profile and in-situ

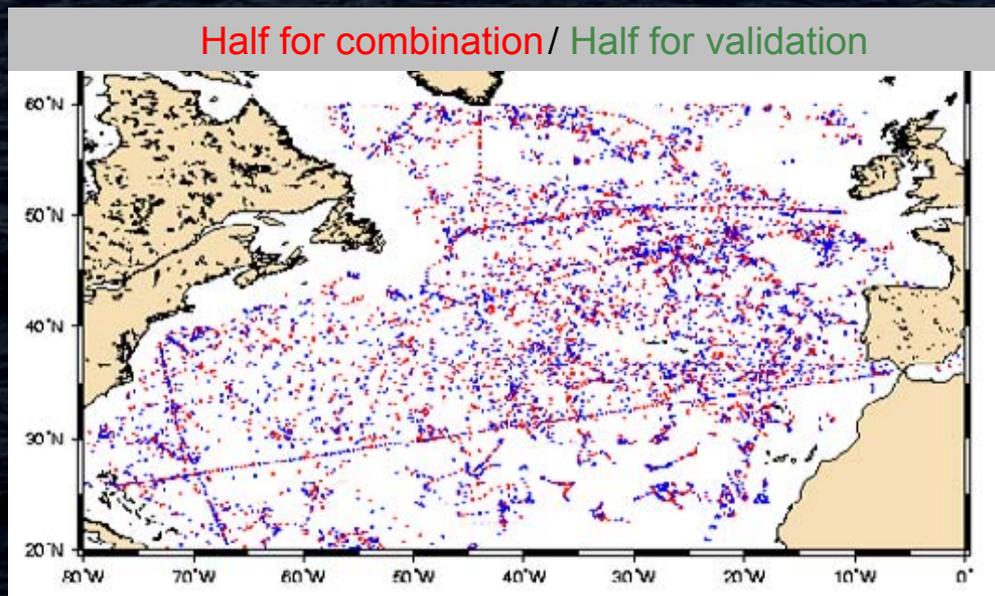




Validation of synthetic and in situ profiles (Le Traon)

Validation with independent data for 2002 year

Half for combination / Half for validation



Methodology

- weekly combination of synthetic and in-situ profiles
- interpolation to the space/time location of independent profiles
- interpolation of independent profiles on the standard Levitus levels
- calculation of the differences between in-situ and combined profiles

- Statistics...to characterize the spatial and vertical structure of the error associated to the 3D combined products



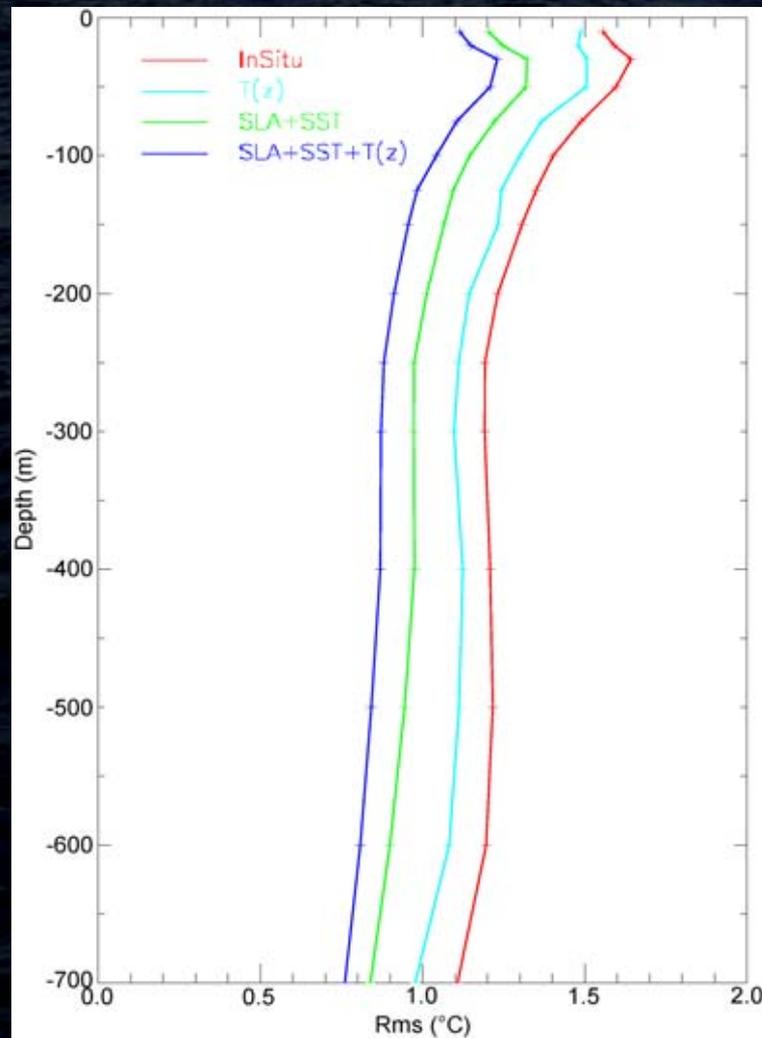
Vertical estimation error (Le Traon)

Satellite data [SLA+SST] + Argo[T(z)]

Combined [SLA+SST+T(z)]

Argo only [T(z)]

Levitus monthly climatology



Mesoscale

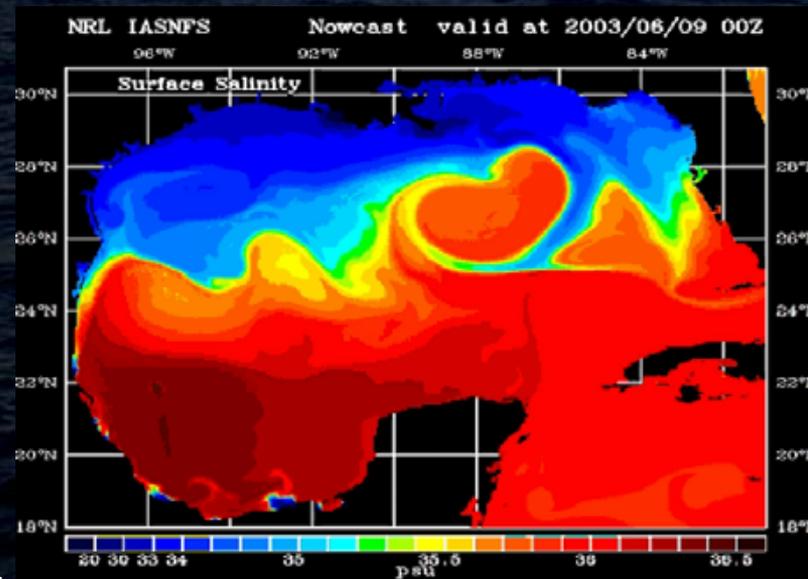
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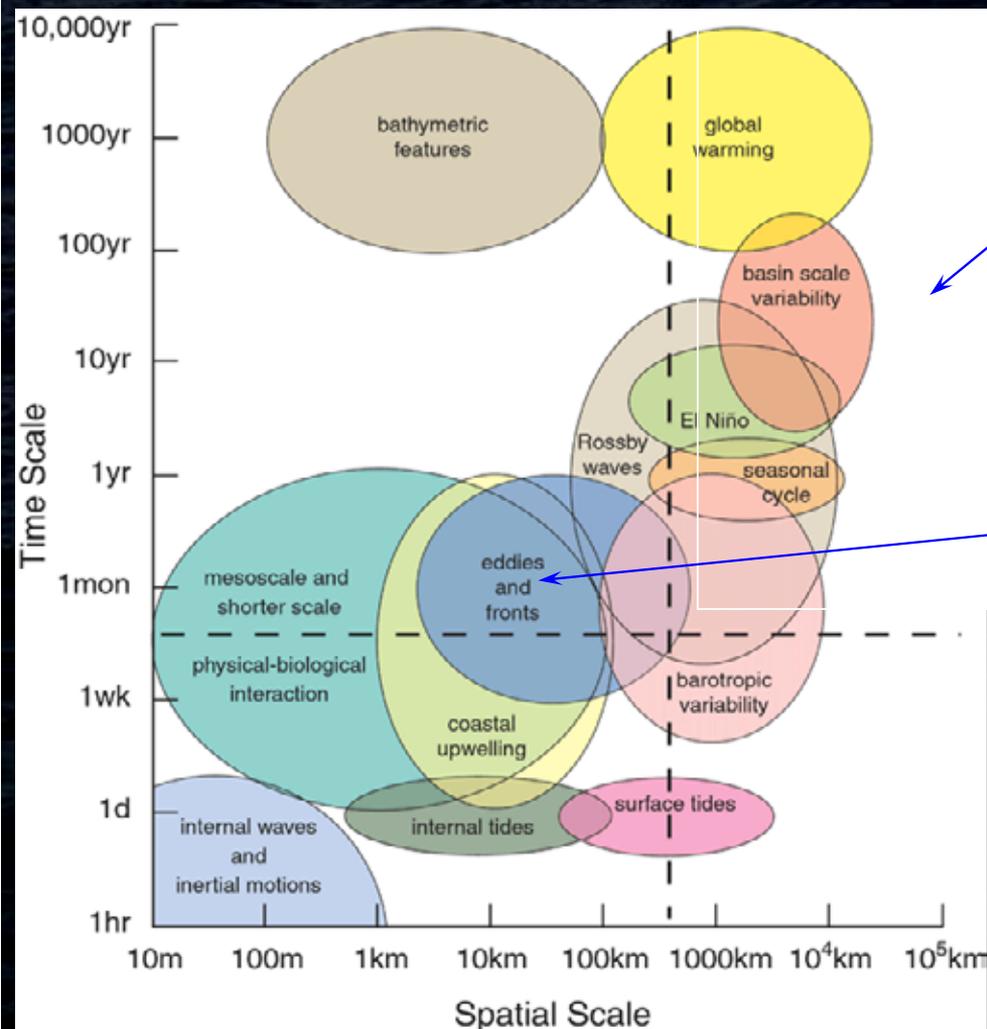
Future potential systems and applications





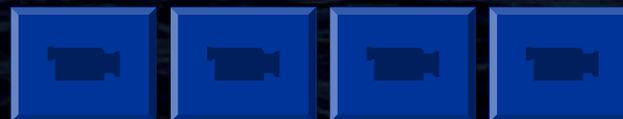
What is resolvable?

Chelton et al., High Resolution Ocean Topography Report, 2001



Resolvable area of Topex/Poseidon or Jason-1

A single nadir altimeter can not resolve the mesoscale.



Bad

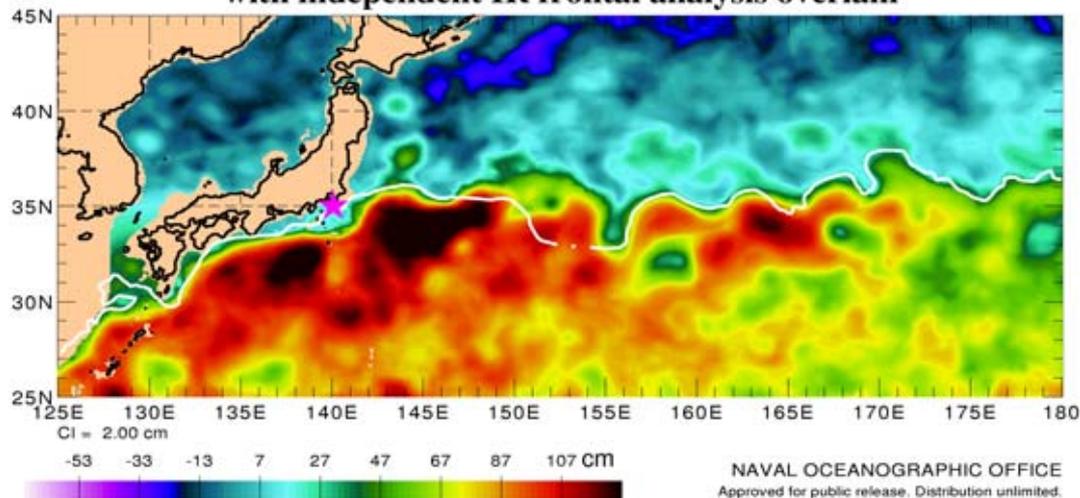
Okay

Best

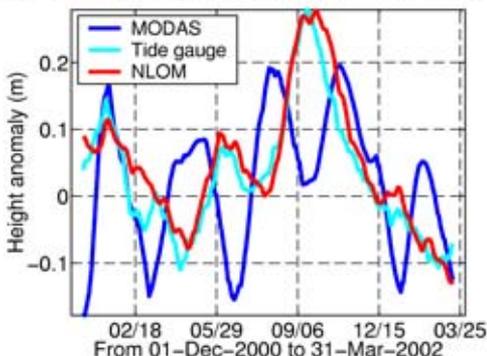


Better Dynamics, Better Results...

1/16° global NLOM SSH analysis for 18 February, 2002
with independent IR frontal analysis overlain



★ Mera, Japan tide gauge comparison



Tide gauge comparison statistics

Mera, Japan 39 concatenated worldwide
28 island, 11 coastal

correlation rmsd correlation rmsd

1/16° NLOM nowcast

.96 3.2 cm .75 5.5 cm

1/8° MODAS model-independent analysis

.34 11.0 cm .65 6.5 cm

30-day running mean tide gauge results

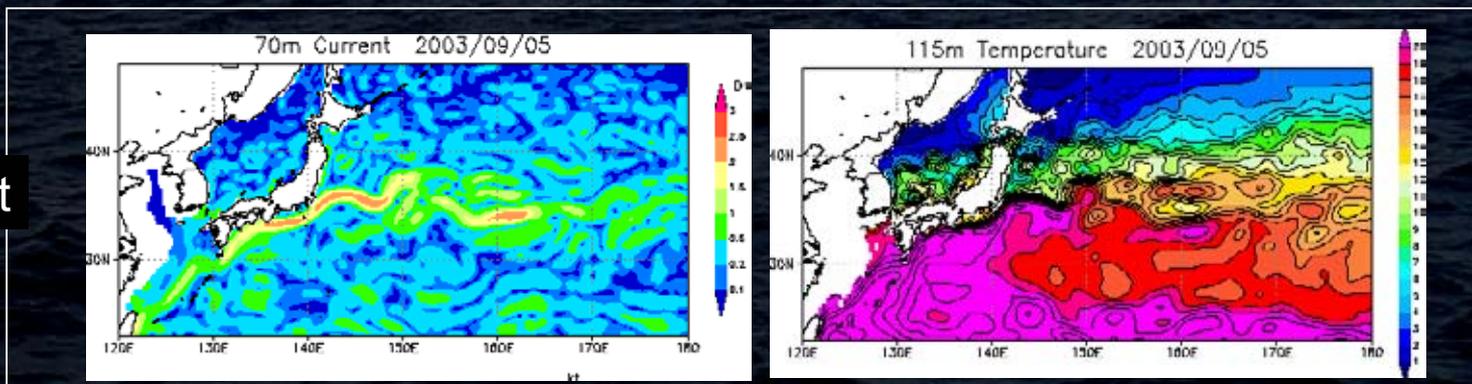
NLOM assimilates altimeter track data using the model as a first guess



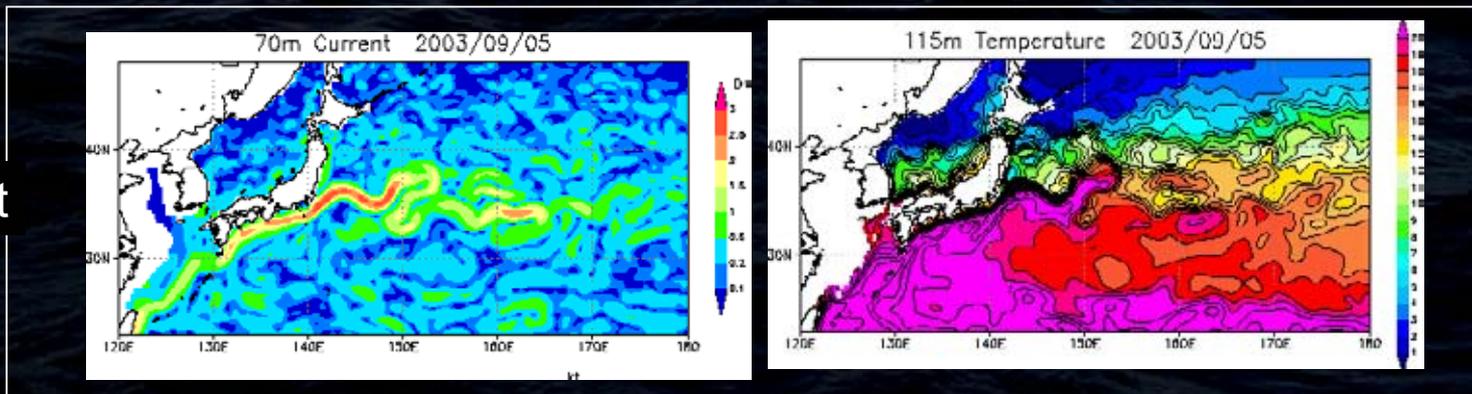
Need for Accurate Model

Japan Met Agency Analysis Cycle (Kuragano)

Forecast



Nowcast



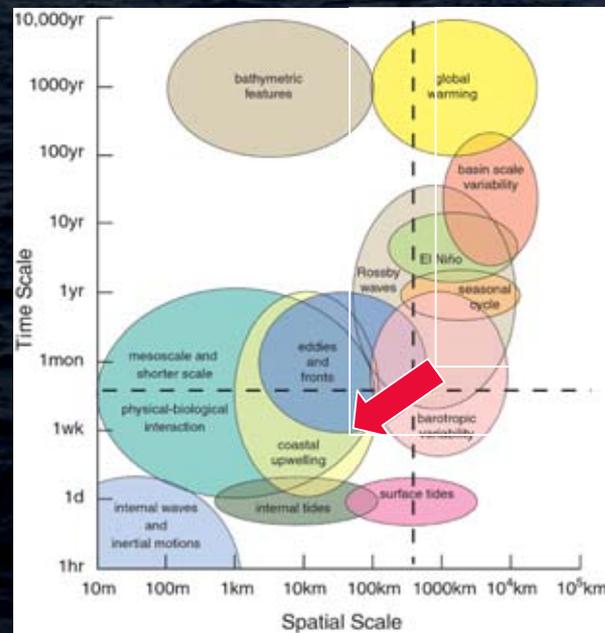
Larger meander at Kuroshio Extension in Nowcast run than that in Assimilation run



But we don't learn anything new

By assimilation (OI, data insertion, EnKF, 4-D Var, ...) we force the solution to match the specified dynamics (to some degree depending on covariance errors).

But we don't learn anything new regarding the dynamics.



To learn anything beyond our present dynamical representations we need

- Independent observations
- Sufficient observations to outweigh our present dynamical understanding
- Assimilation systems to demonstrate we have learned something new

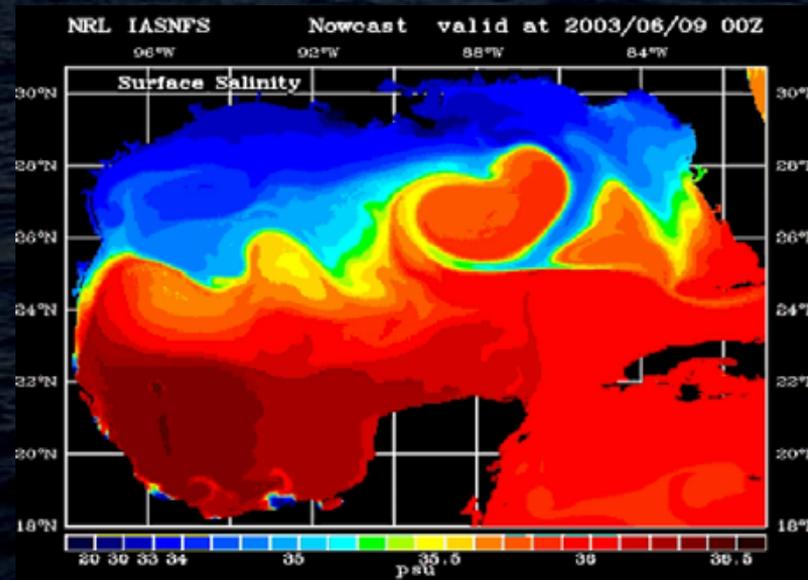
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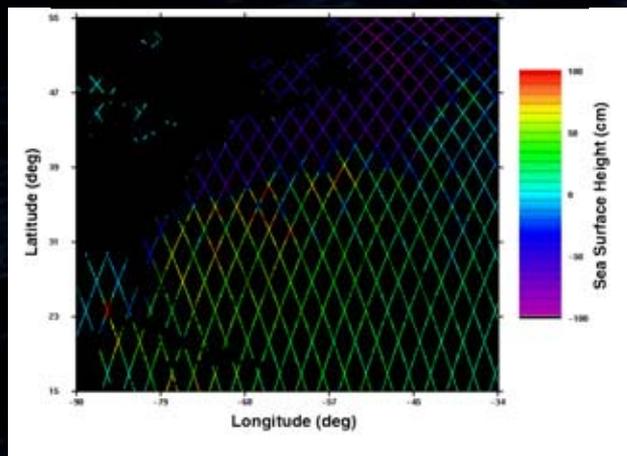


Wide Swath Ocean Altimeter

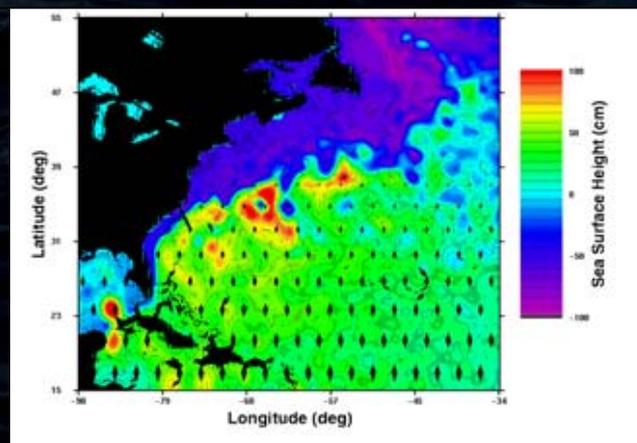
Ocean Surface
Topography Mission



Nadir altimeter

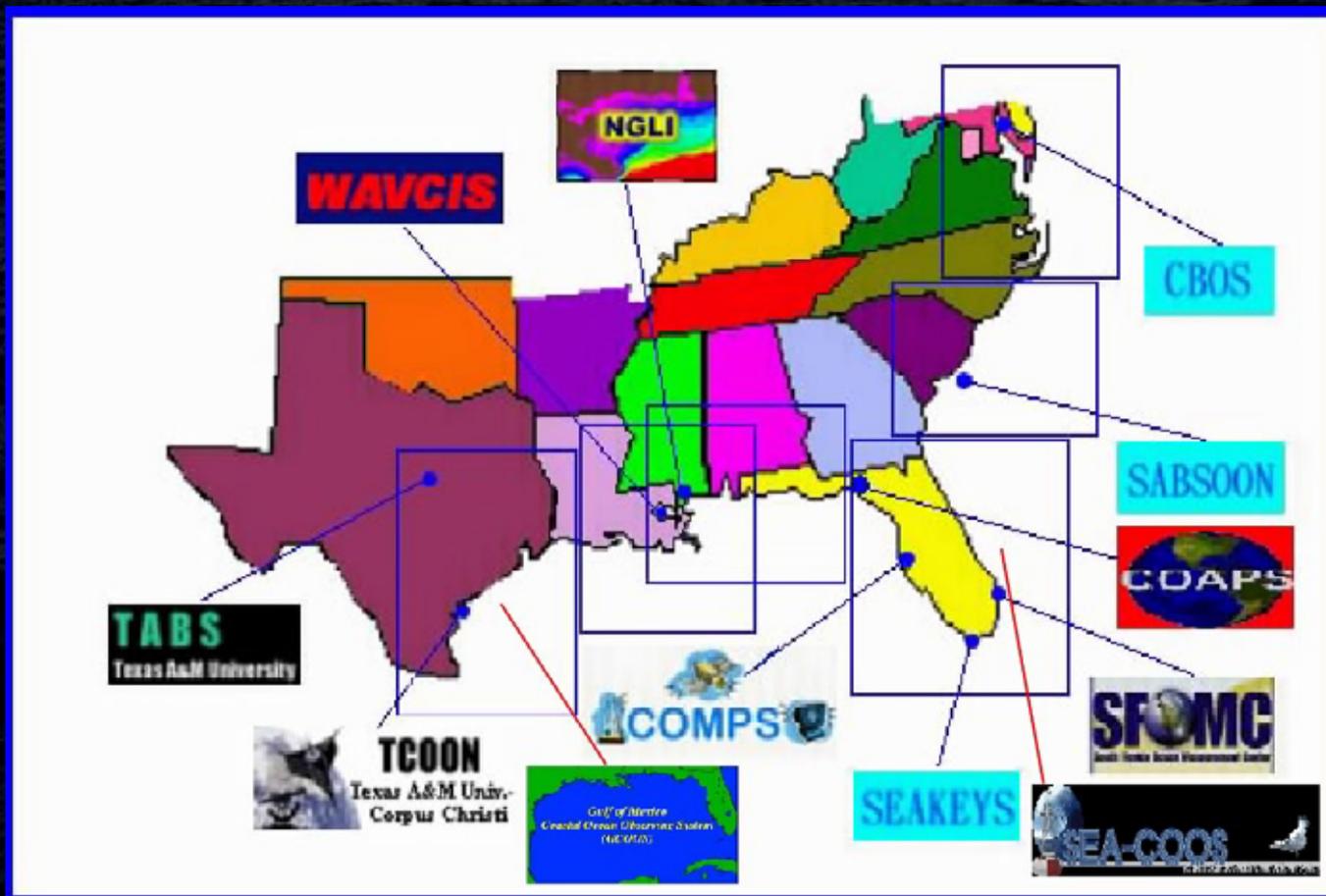


Wide swath altimeter





Integrated Ocean Observation System

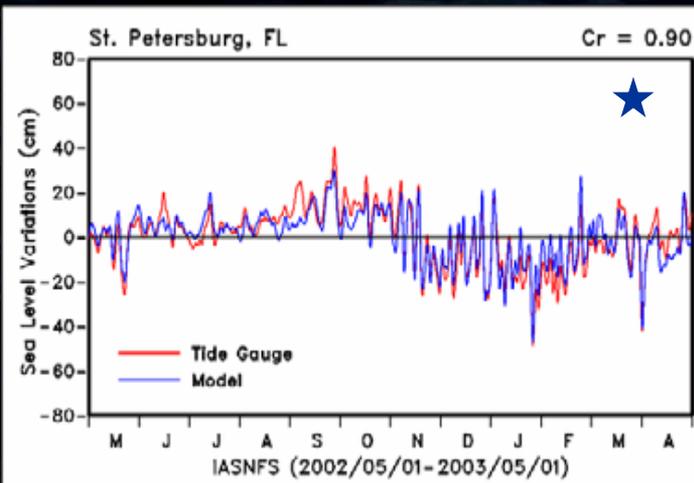
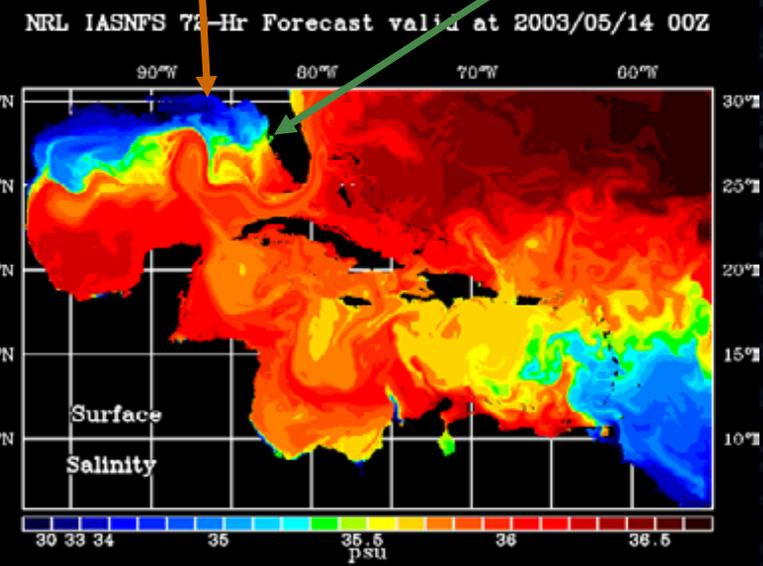
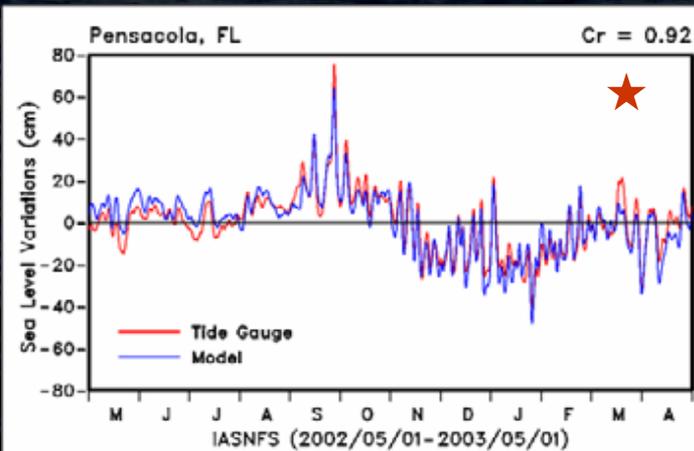
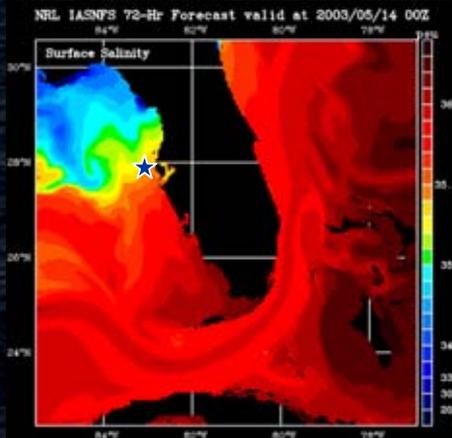
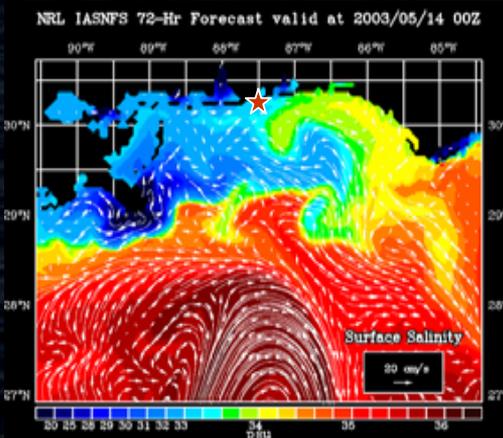


SURA Coastal Ocean Observation System



Regional Modeling

Model/Data Sea Level Comparison





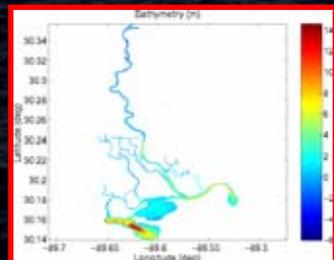
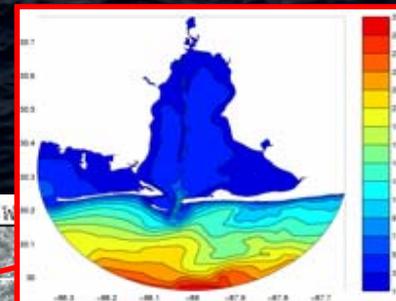
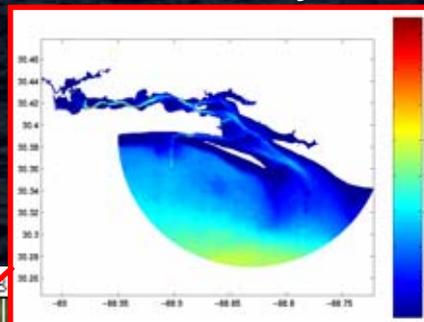
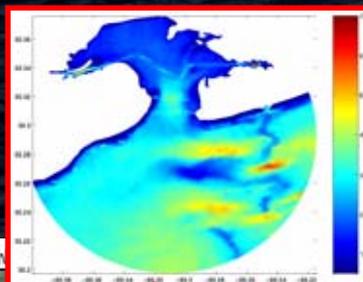
Bay and Estuaries

Bay St. Louis

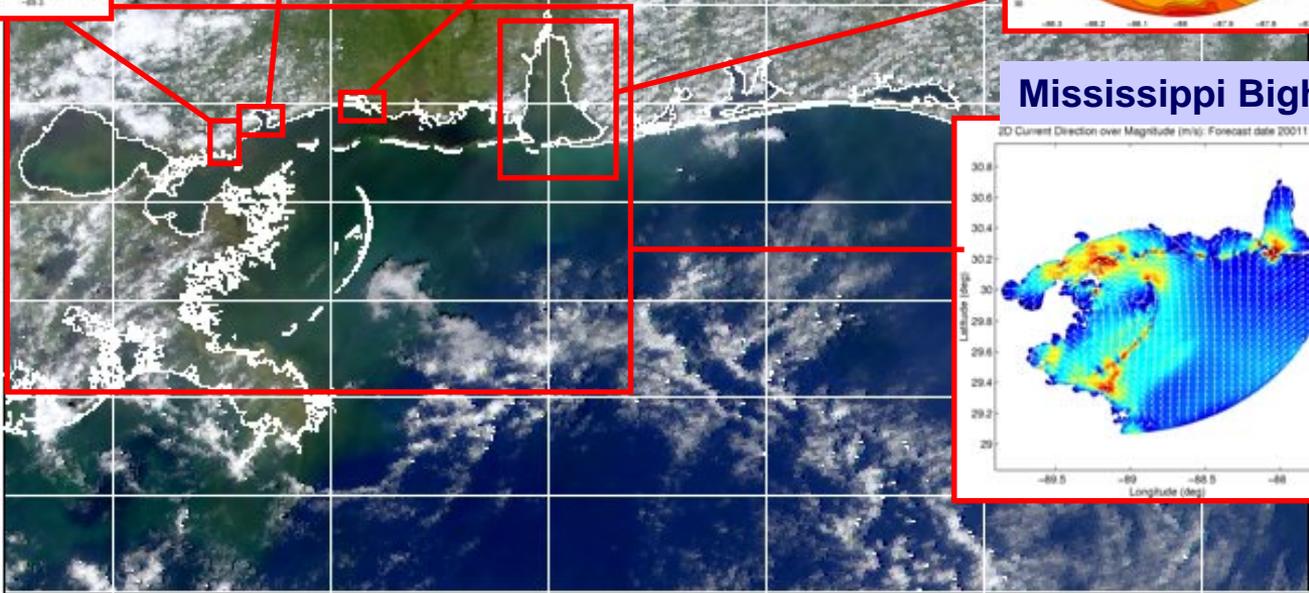
Biloxi Bay

Mobile Bay

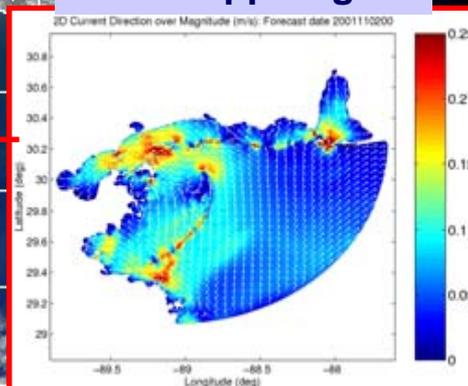
Pearl River



30.40N
30.00N
29.60N
29.20N
28.80N
28.40N

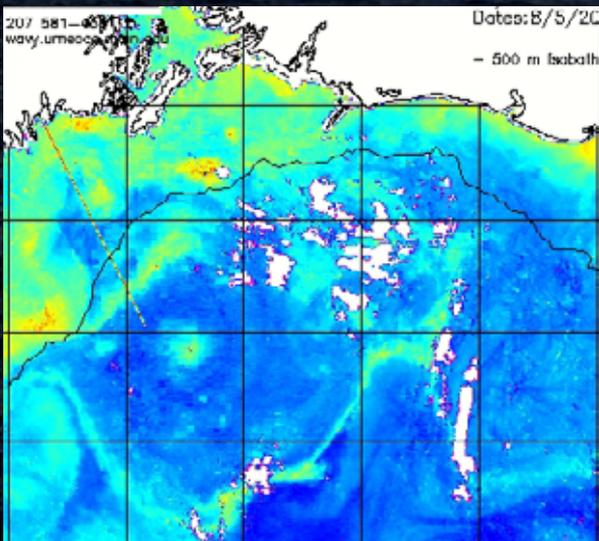
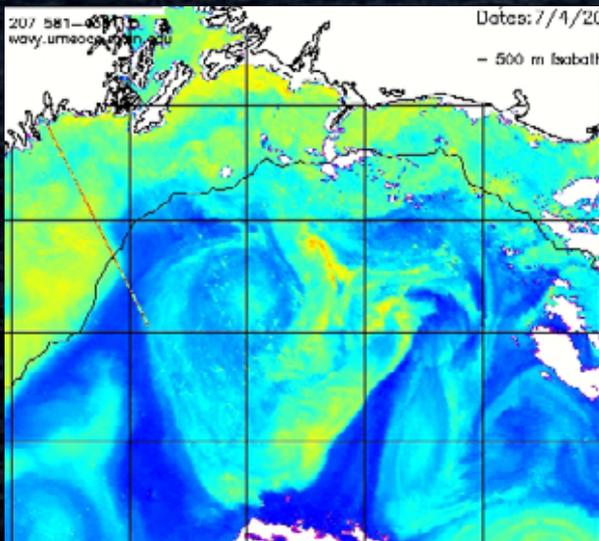


Mississippi Bight

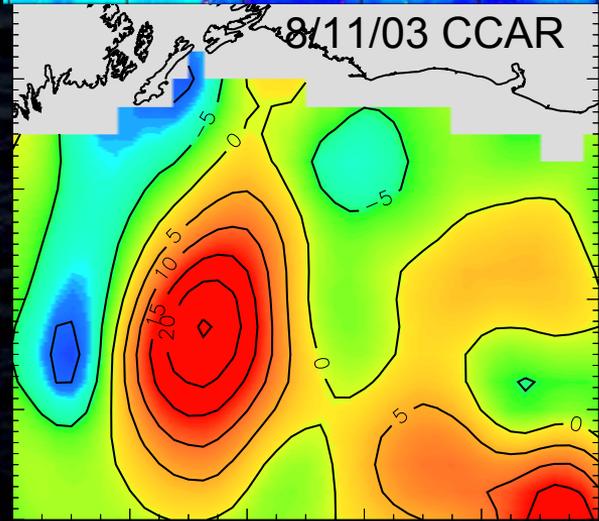
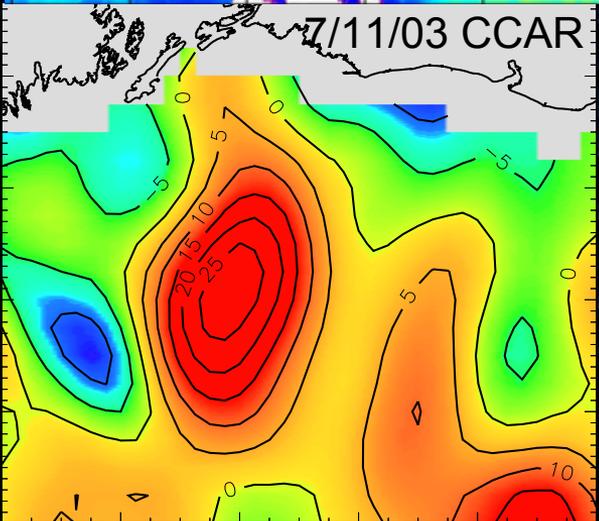




Optical Future (Ted Strub)



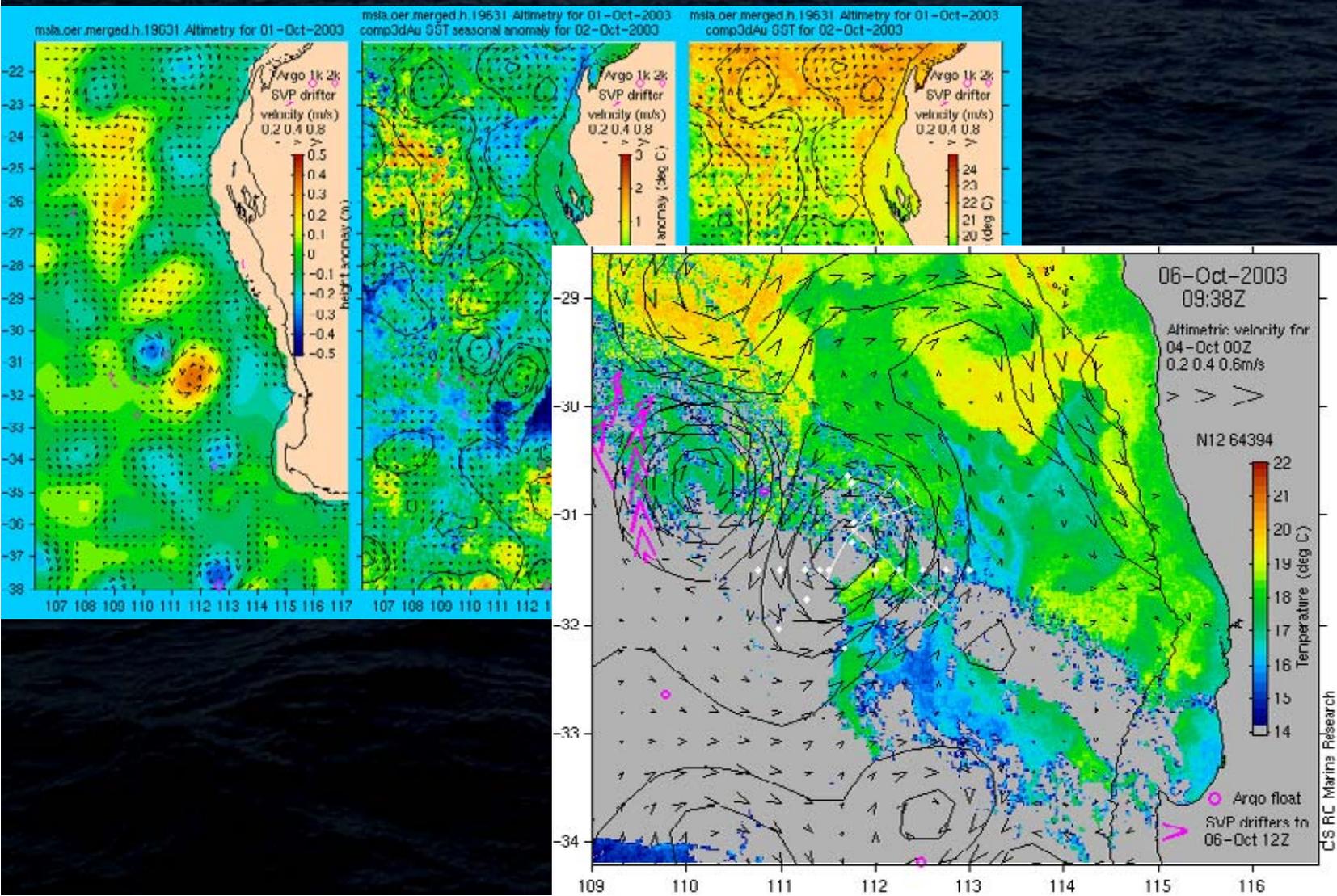
SeaWiFS
From
Andrew
Thomas,
U. Maine



R. Leben



Optical Future (David Griffin)

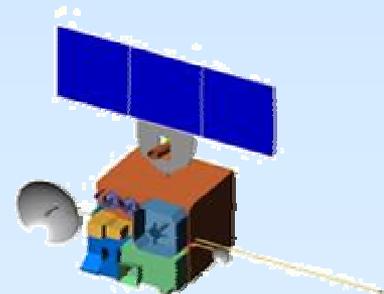




Next Generation Geostationary Satellite - GOES-R



- Responds to national operational environmental requirements
- Provides higher spatial resolution, enhanced spectral information, improved coverage, and more rapid updates to better predict severe weather
 - Better able to discern details that indicate storm formation and provides more accurate wind tracking for model inputs
 - Improved monitoring of volcanic ash, aviation icing hazards, mid-level moisture tracking, and snow/cloud discrimination
 - Able to monitor severe local storms without impacting global coverage
- Program is based on launch readiness by the start of the next decade





GOES R Observational Requirements*

Preliminary Instrument Allocation



Absorbed Shortwave Radiation	Downward Solar Insolation	Rainfall Potential
Aerosol Detection	Dust/Aerosol	Rainfall Rate/QPE
Aerosol Particle Size	Energetic Heavy Ions	Reflected Solar Insolation
Aircraft Icing Threat	Enhanced "V"/Overshooting Top Detection	Sea & Lake Ice/ Displacement and Direction
Atmospheric Vertical Moisture Profile	Fire / Hot Spot Imagery	Sea & Lake Ice/Age
Atmospheric Vertical Temperature Profile	Flood/Standing Water	Sea & Lake Ice/Concentration
Capping Inversion Information	Flood/Standing Water	Sea & Lake Ice/Extent and Characterization
CH ₄ Concentration	Geomagnetic Field	Sea & Lake Ice/Surface Temp
Clear Sky Masks	Hurricane Intensity	Sea Surface Temps
Cloud & Moisture Imagery	Ice Cover/ Landlocked	Snow Cover
Cloud Base Height	Imagery: All-Wx/Day-Nite	Snow Depth
Cloud Ice Water Path	Land Surface (Skin) Temperature	SO ₂ Concentration
Cloud Imagery	Leaf Area Index (LAI)	Solar and Galactic Protons
Cloud Layers / Heights and Thickness	Lightning Detection	Solar Flux: EUV
Cloud Liquid Water	Low Cloud and Fog	Solar Flux: X-Ray
Cloud Optical Depth	Mag Electrons & Protons: Low Energy	Solar Imagery
Cloud Particle Size Distribution	Mag Electrons & Protons: Med & High Energy	Surface Albedo
Cloud Phase	Microburst Winds	Surface Emissivity
Cloud Top Height	Moisture Flux	Suspended Matter
Cloud Top Pressure	Ocean Currents	Total Precipitable Water
Cloud Top Temperature	Ocean Color	Total Water Content
Cloud Type	Ocean Optical Properties	Turbulence
CO Concentration	Ocean Turbidity	Upward Longwave Radiation
CO ₂ Concentration	Oil Spill Location	Vegetation Fraction
Convection Initiation	Ozone Layers	Vegetation Index
Derived Motion Winds	Ozone Total	Visibility
Derived Stability Indices	Pressure Profile	Volcanic Ash
Downward Longwave Radiation	Probability of Rainfall	
	Radiances	

ABI – Advanced Baseline Imager

HES – Hyperspectral Environmental Suite

SEM – Space Environment Monitor

SXI – Solar X-Ray Imager

GLM – GOES Lightning Mapper

* Does not reflect individual geographic coverage requirements.



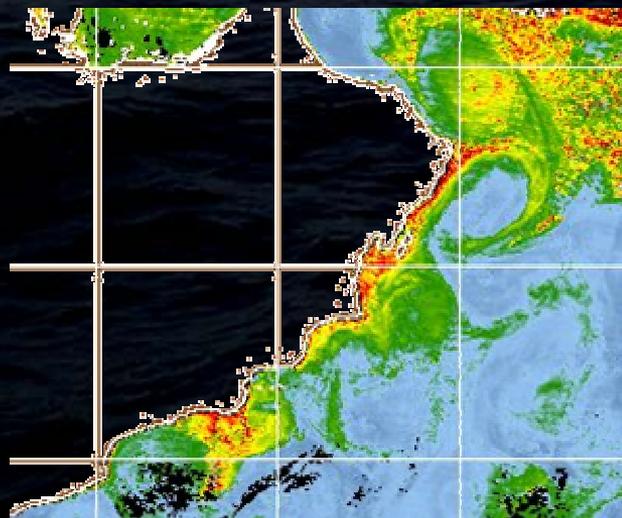
Optical Future (John Kindle)

Assimilation of Satellite Imagery and Numerical Models

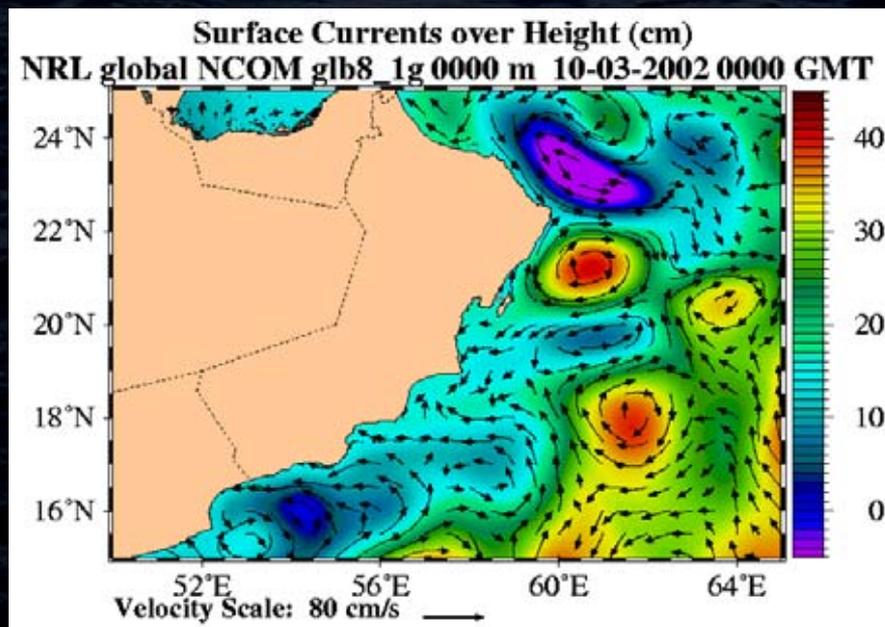


SeaWiFS composite

Sep 30-Oct 7 2002



Global NCOM: Oct 3 2002

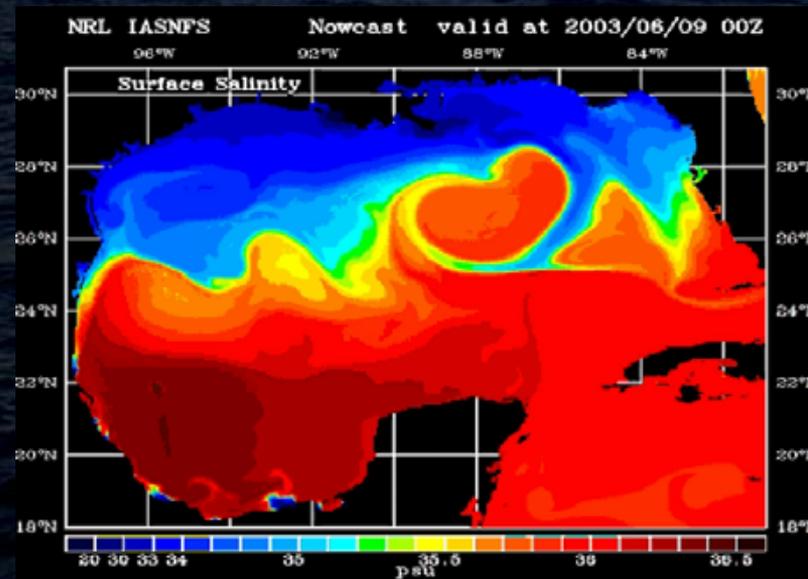


Conclude

Altimeter observations are essential for understanding the synoptic mesoscale dynamics and fluxes

Mesoscale interacts with coastal region and produces observable effects in non-dynamic observables

Altimeter community, biology, geology, chemistry, ... communities need to bring the components together





End
