

Observation of swell pattern and energetic surface currents with synthetic aperture radar

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retrieval of Surface Currents from Doppler shift



Residual velocities





Comparison with a lagrangian drifter (15m depth)



SAR observation of strong surface currents



Map of surface velocity of the Agulhas Current using SAR Doppler information



Altimetry derived geostrophic surface current : 3 days mean

SAR vs. conventional altimetric surface currents

Projection on SAR radial direction





Radial Sea surface Current from Altimetry 2007-09-16



SAR vs. conventional altimetric surface currents

Radial Doppler Velocity 2007-09-16 ASAR & ALTIMETER 2007-09-16 - ALTIMETER e- ASAR -34 34°S -35 -36 3608 <u></u> -37 latitude [latitude [°] 38⁰S -39 -4(40°S -A · -42 42°S radial velocity [m/s] 22°E 24°E 26⁰E 28⁰E 30⁰E 32⁰E longitude [°] Mesoscale eddie 0.5 -2 -1.5 - 1 -0.5 0 1.5 m/s

Agulhas main stream

Challenge in coastal altimetry : what percentage of surface current magnitude signs in ocean topography

Mean current

Annual mean ascending radial component of surface current over Agulhas and Gulf stream





Meridional mean surface current

SAR vs.



MDT Rio09

CLS

















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CLS esa UNCLASSIFIED: 1/32° Global NLOM CURRENT/SPEED LAYER 1 ANALYSIS: 20100606 90W 85W 0.50m/s 64 NAVAL OCEANOGRAPHIC OFFIC Approved for public release. Distribution unlimite







SWORD SWath Ocean Radar Doppler



CLS

SWORD SWath Ocean Radar Doppler



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Wave transformation in current

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 Significant swell height and dominant direction



SAR swell products



- Only SAR enables observation of spatial variations of 2D swell spectrum measurement
 <u>http://soprano.cls.fr/waveProducts/</u>
- Available information of the swell component
 - Significant wave height
 - Wavelength (size of arrow)
 - Dominant direction
- General breakthrough
 - First high resolution wave field observation at 5km resolution
 - Wave inversion in shallow water

Observation of Swell variation in coastal area



<u>Swell nonlinearities in shallow water</u>



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Infragravity waves and Sea level remote sensing

Generalities on infragravity (IG) waves

1. Unlike gravity waves, free IG waves are typically longer than altimeter footprints

2. Infragravity (IG) waves are generated as « bound waves » by the nonlinear difference interaction of « normal waves » of frequencies f1 and f2 and wavenumbers k1 and k2:

 $f_IG = f1 - f2$ (0.001 to 0.05 Hz)

and $k_IG = k1 - k2$



At this stage the wavelength of IG waves is the lenght of wave groups (0 (1 km)). Forced IGs are strongly amplified at the coast (amplitude can reach 1 m on the beach face).

As the forcing wave groups are dissipated, the IG waves are realeased as free modes. This energy is mostly trapped in the coastal zone, but a fraction of that energy leaks into the deep ocean. Free waves have the same frequency (0.001 to 0.05 Hz) ... but the wavelength is given by the linear surface gravity wave dispersion: assuming shallow water: $L_IG = sqrt(g^*D)/f_IG$ [more generally, f^2 = g k tanh(kD)/(2*pi)^2] For D= 4000 m and a typical f_IG=0.02 Hz, this gives L_IG = 9 km

Conclusion



- 1cm at 10km resolution shall necessitate validation with improved MDT and instantaneous surface current field at finer resolution
- Strong heterogeneity of wave field in dynamical area. Needs to estimate the impact on altimeter interferometric measurement precision
- Wave bends as function of vorticity (and the wave group velocity) : wave impact and high resolution topography anomalies like SSB are not independent of surface currents.
- SAR can provide an independant high resolution NRCS and Doppler shift map : when defining the orbit, having SWOT swath inside an existing SAR swath would bring interesting comparisons opportunities.