October 21, 2010

# Two science questions that ca only be addressed with SWO

Raffaele Ferrari Massachusetts Institute of Technology in collaboration with bernathey (MIT), A. Naveira Garabato (SOC), M. Nikura



# Lateral eddy transport Vertical eddy transport

Eddies and winds in the Southern Ocean (SO) set the global ocean tratification below the surface thermoclines (Wolfe and Cessi, '10)



# Slope of isopycnals in SO = $\frac{\tau_{SO}}{fK_{SO}}$

Eddies and winds in the Southern Ocean set the strength of the neridional overturning circulation (Wolfe and Cessi, '10; Nikurashin and Vallis, '10)



- nds over the Southern Ocean have strengthened over the last 30 ye ompson and Solomon, '02)
- s unclear whether ocean CO<sub>2</sub> uptake will increase or decrease in oonse to wind increase (Le Quéré et al., '07; Meredith and Hogg, '06) take depends on how *Kso* changes in response to wind change



ere are spatial variations in  $K_{SO}$ 

*Kso* is suppressed in the core of the Antarctic Circumpolar Current *Kso* is enhanced at depth below the core of the ACC

$$C_{SO} = 0.32 \frac{g}{f} \frac{\sqrt{\eta'^2}}{1+8|\nabla \overline{\eta}|^2/|\nabla \eta'|^2} = 0.32 \frac{g}{f} \frac{\sqrt{\eta'^2}}{1+8U^2/EKT}$$



Marshall et al., '04

# ppression of *Kso* is confined to ACC frontal jets ontal jets

### Drake Passage Section (WOCE) from Garabato and Ferrari, '10



aditional altimeters have a resolution of O(100)km

- ACC fronts cannot be resolved
- VOT will have a resolution of O(10)km
- ACC fronts will be resolved

magnitude and structure of  $K_{SO}$  will be measured for the first time





# Lateral eddy transport Vertical eddy transport

change of tracers and nutrients between the ocean surface and inte centrated at fronts

onts are speculated to regulate primary productivity in the ocean

### Lateral dispersal of tracer - due to the slumping of a ML front



- rtical transport at fronts in the QG approximation is determined by norizontal density gradients
- norizontal velocity shears
- th quantities are dominated by frontal scales of O(1-50)km
- peyre and collaborators suggested to use SWOT and QG theory to i ical transport of tracers from sea surface height

$$f^2 \psi_{zz} + N^2 \psi_{yy} = -2\nu_y b_y$$



# Are QG estimates of vertical transport accurate?

- nulations of baroclinically unstable fronts are run with a QG model a nitive Equation model
- e goal is to test whether a QG model captures accurately the vertica sport of tracers





$$b = N^2 z + rac{M^2}{\ell_0} \cos \ell_0 y c$$
  
 $u = rac{M^2}{fm_0} \sin \ell_0 y \sin m_0 z$   
 $Ro = rac{M^2}{Nf}$   
 $Ri = Ro^{-2}$ 

# Are QG estimates of vertical transport accurate?

- nulations of baroclinically unstable fronts are run with a QG model a nitive Equation model
- e goal is to test whether a QG model captures accurately the vertica sport of tracers



Ro simulations are QG-like ep horizontal KE spectrum (k<sup>-3</sup>)

 $\frac{w^2}{2 + v^2} \ll 1$ Ro=0.13, Ri=64



High Ro simulations are not QGflat horizontal KE spectrum (k<sup>-5</sup>/<sub>2</sub>)

•  $\frac{w^2}{u^2 + v^2}$  approaches 1

Ro=0.42, Ri=5.7



rtical velocity is underestimated by QG approximation at Ro=O(1) rtical tracer transport of tracers is well captured by QG approximatio n for Ro=O(1)



Full resolution of lateral eddy diffusivity Partial resolution of vertical transport at fronts A new window on ocean productivity



#### Matthew Mazloff (Ph.D. thesis)

78° South to 24.7° South
1/6° Horizontal resolution;
42 depth levels (partial cells)
similar setup to ECCO-GODAE



## Kinetic Energy



## Horizontal SeaSoar section

## KE and PE horizontal spe



# Equipartition of KE and PE Spectral rolloff of k-2



### Horizontal SeaSoar section

## Rossby number





# Equipartition of KE and PE

Spectral rolloff of k-2

Eddies and winds in the Southern Ocean set the strength of the neridional overturning circulation (Nikurashin and Vallis, '10)



 $\kappa L$ 117 12/2