# Dynamical separation of stationary and non-stationary internal tides

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June 2019



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- Present dynamical equations for the stationary tide
- Show that meanflow effects explain the decay of the stationary tide
- Parameterize meanflow effects with an eddy diffusivity, and apply the parameterization to the global ocean

Substitute 
$$H\mathbf{u}'(\mathbf{x}, z, t) = \sum_{n=1}^{\infty} \mathbf{U}_n(\mathbf{x}, t)\phi_n(z)$$
 and  $p'(\mathbf{x}, z, t) = \sum_{n=1}^{\infty} p_n(\mathbf{x}, t)\phi_n(z)$ 

#### Horizontal dependence (Shallow water equations)

$$\mathbf{U}_{nt} + f\mathbf{k} \times \mathbf{U}_{n} = -H\nabla p_{n}$$

$$\frac{Hp_{nt}}{c_{n}^{2}} = -\nabla \cdot \mathbf{U}_{n}$$

#### Vertical dependence (a time-independent eigenvalue problem)

$$\frac{\partial}{\partial z} \left( \frac{1}{N^2} \frac{\partial \phi_n}{\partial z} \right) + \frac{1}{c_n^2} \phi_n = 0$$

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#### Variable topography (Coupled shallow water equations)

$$\mathbf{U}_{nt} + f\mathbf{k} \times \mathbf{U}_{n} = -H\nabla p_{n} - \sum_{m=0}^{\infty} p_{m} \mathbf{T}_{mn}$$

$$\frac{Hp_{nt}}{c_{n}^{2}} = -\nabla \cdot \mathbf{U}_{n} + \sum_{m=0}^{\infty} \mathbf{U}_{m} \cdot \mathbf{T}_{mn}$$

#### Topographic coupling coefficients (where H and $\phi$ vary with ${f x})$

$$\mathbf{T}_{mn} = \frac{1}{H} \int_{-H}^{0} \phi_n \nabla \phi_m \mathrm{d}z$$

#### Leading-order meanflow interaction

$$\mathbf{U}_{nt} + \sum_{m=0}^{\infty} \nabla \cdot \left( \overline{\mathbf{u}}_{mn}^{\mathsf{T}} \mathbf{U}_{m} \right) + f \mathbf{k} \times \mathbf{U}_{n} = -H \nabla p_{n} - \sum_{m=1}^{\infty} p_{m} \mathbf{T}_{mn}$$

$$\frac{H p_{nt}}{c_{n}^{2}} + \sum_{m=0}^{\infty} \nabla \cdot \left( \frac{\overline{\mathbf{u}}_{mn} H p_{m}}{c_{n}^{2}} \right) + \frac{\delta c_{n}^{2}}{c_{n}^{2}} \nabla \cdot \mathbf{U}_{n} = -\nabla \cdot \mathbf{U}_{n} + \sum_{m=0}^{\infty} \mathbf{U}_{m} \cdot \mathbf{T}_{mn}$$

#### Meanflow coupling coefficients

$$\begin{aligned} & \overline{\mathbf{u}}_{mn}(\mathbf{x},t) = \frac{1}{H} \int_{-H}^{0} \overline{\mathbf{u}}(\mathbf{x},z,t) \phi_{m} \phi_{n} \mathrm{d}z \\ & \delta c_{n}^{2}(\mathbf{x},t) = \frac{1}{H} \int_{-H}^{0} \delta N^{2}(\mathbf{x},z,t) \Phi_{n} \Phi_{n} \mathrm{d}z \end{aligned}$$

#### Leading-order meanflow interaction

$$\mathbf{U}_{nt} + \sum_{m=0}^{\infty} \nabla \cdot \left( \overline{\mathbf{u}}_{mn}^{\mathsf{T}} \mathbf{U}_{m} \right) + f \mathbf{k} \times \mathbf{U}_{n} = -H \nabla p_{n} - \sum_{m=1}^{\infty} p_{m} \mathbf{T}_{mn}$$

$$\frac{H p_{nt}}{c_{n}^{2}} + \sum_{m=0}^{\infty} \nabla \cdot \left( \frac{\overline{\mathbf{u}}_{mn} H p_{m}}{c_{n}^{2}} \right) + \frac{\delta c_{n}^{2}}{c_{n}^{2}} \nabla \cdot \mathbf{U}_{n} = -\nabla \cdot \mathbf{U}_{n} + \sum_{m=0}^{\infty} \mathbf{U}_{m} \cdot \mathbf{T}_{mn}$$

#### **Assumptions**

- Small Rossby and Froude number  $(\epsilon = U/fL \ll 1)^*$
- Geometric approximation (for dynamical stability)
- Simultaneous meanflow and topographic effects are weak

<sup>\*</sup>See asymptotic derivation by Wagner, Ferrando, and Young (2017)

# Solving the system

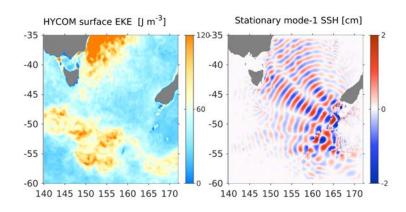
#### Coupled shallow water model\* (CSW)

- Finite differences on a spherical C grid
- Adams-Bashforth time-stepping algorithm
- Damped by linear/quadratic drag, viscosity, or sponge
- Forced by prescribed surface tide velocities
- C code without meanflow available at Bitbucket.org
- Matlab code with meanflow available by email

Resolution	# modes	cores	RAM [GB]	speed [cycles/hr]
1/25°	4	16	30	10
$1/50^{\circ}$	4	128	150	11
$1/100^{\circ}$	4	256	750	2.3

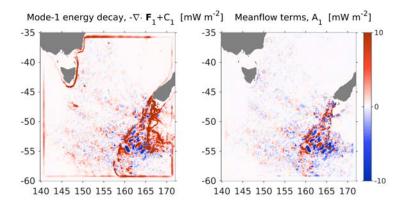
<sup>\*</sup>The model is described in Kelly et al. (2016) and Griffiths and Grimshaw (2007)

#### Simulation of Tasman Sea for 2015



- TPXO M2 surface tides
- Smith and Sandwell bathymetry
- HYCOM meanflow
- Horizontal viscosity  $\nu_T = 27.5 \text{ m}^2 \text{ s}^{-1}$  (for stability)

## Do meanflow effects dissipate the mode-1 tide?...No



 Lateral sponges dissipate most mode-1 energy (meanflow terms nearly average to zero)

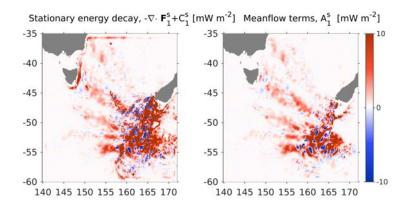
# Stationary tide equations

$$\begin{aligned} \mathbf{U}_{nt}^{s} + f\mathbf{k} \times \mathbf{U}_{n}^{s} &= -H\nabla p_{n}^{s} - \sum_{m=1}^{\infty} p_{m}^{s} \mathbf{T}_{mn} - \left[ \sum_{m=0}^{\infty} \nabla \cdot \left( \overline{\mathbf{u}}_{mn}^{T} \mathbf{U}_{m} \right) \right]^{s} \\ \frac{Hp_{nt}^{s}}{c_{n}^{2}} &= -\nabla \cdot \mathbf{U}_{n}^{s} + \sum_{m=0}^{\infty} \mathbf{U}_{m}^{s} \cdot \mathbf{T}_{mn} - \left[ \sum_{m=0}^{\infty} \nabla \cdot \left( \frac{\overline{\mathbf{u}}_{p,mn} Hp_{m}}{c_{m}c_{n}} \right) + \frac{\delta c_{n}^{2}}{c_{n}^{2}} \nabla \cdot \mathbf{U}_{n} \right]^{s} \end{aligned}$$

#### Notes:

- Stationary variables (e.g.,  $\mathbf{U}_n^s$  and  $p_n^s$ ) are harmonic fits or ensemble averages
- Signals are orthogonal with respect to time averaging:  $\langle \mathbf{U}_n^s p_n \rangle = \langle \mathbf{U}_n^s p_n^s \rangle$  and  $\langle \mathbf{U}_n^s (p_n p_n^s) \rangle = 0$ .
- Meanflow terms (square brackets) are fixed, but unclosed (depend on non-stationary tide).

#### Do meanflow effects dissipate the stationary mode-1 tide?...Yes



Meanflow terms explain most stationary tide decay

 $A_1$  is also the generation map for non-stationary tides

# Parameterizing meanflow effects on the stationary tide

#### Greatly simply the model with an eddy viscosity

$$-\left[\sum_{m=0}^{\infty} \nabla \cdot \left(\overline{\mathbf{u}}_{mn}^{T} \mathbf{U}_{m}\right)\right]^{s} \approx \nu_{T} \nabla^{2} \mathbf{U}_{n}^{s}$$

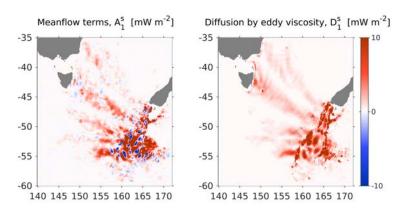
where\*

$$\nu_T = \Gamma L \sqrt{2 \text{EKE}}$$

- EKE is observed eddy kinetic energy
- Γ is the mixing efficiency (a free constant)
- L is the mixing length (Rossby radius or eddy diameter)

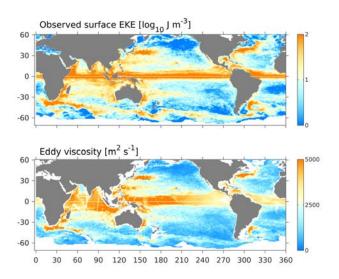
<sup>\*</sup>See Klocker and Abernathy (2014) for details on  $\nu_T$ .

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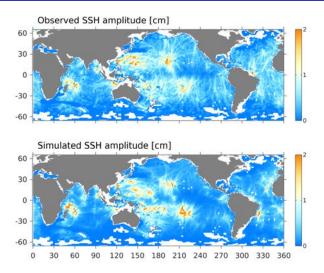


• 47% of  $A_1$  variance explained when averaged to  $1^\circ$ 

# Global parameterization of meanflow effects



# Global parameterization of meanflow effects



 $\bullet$  68% of M2 amplitude variance explained when averaged to  $1^{\circ}$ 

HRET SSH courtesy of Ed Zaron

## Summary

#### Conclusions

- Meanflow effects explain stationary mode-1 tide decay
- Meanflow effects can be parameterized by eddy viscosity

#### Additional thoughts

- Determine the stationary and non-stationary tide separately?
- 2 Still no obvious dominant source of final mode-1 dissipation

**Data sources:** Jim Richman & Jay Shriver (HYCOM), volkov.oce.orst.edu (TPXO), topex.ucsd.edu (Smith & Sandwell bathy.), Ed Zaron (HRET M2 tides), Bo Qiu (AVISO EKE)